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GEOTHERMAL RESOURCE ANALYSIS IN TWIN FALLS COUNTY, IDAHO

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GEOTHERMAL RESOURCE ANALYSIS IN TWIN FALLS COUNTY, IDAHO

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TABLE OF CONTENTS

Introduction
Purpose and Approach3
Previous Work
Geologic Studies3
Geothermal Studies4
Climate5
Physiography5
Geologic Framework7
Stratigraphy8
General Statement8
Pre-Cenozoic Rocks (MPz)8
Tertiary Idavada Volcanic Group9
Idavada Ash-Flow Tuffs in the Cassia Mountains (Tiv)10
Older Rhyolite (Not Exposed)
Lacustrine Sediments11
Subsurface Alteration and Mineralization
Shoshone Falls Rhyolite12
Tertiary and Quaternary Basalts (Tb, QTb and Qb)14
Geochemistry of Idavada Rhyolites15
Tectonics and Structure16
Hydrology of Thermal Systems20
Aquifer Characteristics20
Aquifer Testing22
Water Chemistry23
Monitoring of Thermal Wells23
Annual Discharge36
Preliminary Conceptual Model of the Thermal System37
Summary and Recommendations for Further Studies40
References

LIST OF FIGURES

Figure 1	Index Map2							
Figure 2	Tectonics and Structure18							
Figure 3	Generalized Stratigraphy of the Study Area21							
Figure 4	Well # 1 Hydrograph25							
Figure 5	Well # 2 Hydrograph26							
Figure 6	Well # 3 Hydrograph27							
Figure 7	Well # 4 Hydrograph28							
Figure 8	Well # 5 Hydrograph29							
Figure 9 .	Well # 7 Hydrograph30							
Figure 10	Well # 11 Hydrograph31							
Figure 11	Well # 12 Hydrograph32							
Figure 12	Well # 13 Hydrograph33							
Figure 13	Conceptual Model of the Twin Falls - Banbury							
	Thermal System39							
	Tables							
Table 1	Well Information and Frequency of Measurements24							
APPENDIX								
Appendix A	Driller's Logs							
Appendix B	Generalized Stratigraphic Section for the Cassia							
	Mountains							
Appendix C	Chemical Analysis of Idavada Pyroclastics and the							
	Shoshone Falls Rhyolite Techniques and Results							

Plate 1

Geologic Map

ABSTRACT

Increased utilization of the geothermal resource in the Twin Falls - Banbury area of southern Idaho has resulted in noticeable declines in the artesian head of the system. In order to determine the nature of the declines, a network of wells was identified for monitoring shut-in pressures and temperatures. In addition, a compilation of data and reconnaissance of the areal geology was undertaken in order to better understand the geologic framework and its relationship to the occurrence of the thermal waters in the system.

The geothermal resource of the Twin Falls - Banbury system is characterized by temperatures between 30° and 70°C (86° to 158°F) and shut-in well pressures of 14 to 250 psi. The thermal water occurs in rhyolitic ash-flow tuffs and lava flows of the Tertiary Idavada Volcanic Group. Permeability of the reservoir rocks results from tectonic and cooling fractures, intergranular porosity of the non-welded tuffs and voids left between successive flows. The system is recharged by rain and snow falling on the Cassia Mountains to the south. Northward dipping volcanic strata channel the water toward the center of the Snake River Plain and into northwest trending structure zones which cross the area from Hollister to Banbury Hot Springs. The heat source is thought to be the regionally high heat flow supplemented in parts of the system by deep circulation in structure zones.

The results of the monitoring indicate that while water temperatures have remained constant, the system shows a gradual overall decline in artesian pressure superimposed on fluctuations caused by seasonal use of the system. Well testing and the similarity of hydrographs resulting from well monitoring throughout the area suggest that there are no major hydrologic barriers to thermal water movement in the system and that wells are affected by increases and decreases in utilization of nearby wells.

DISCLAIMER

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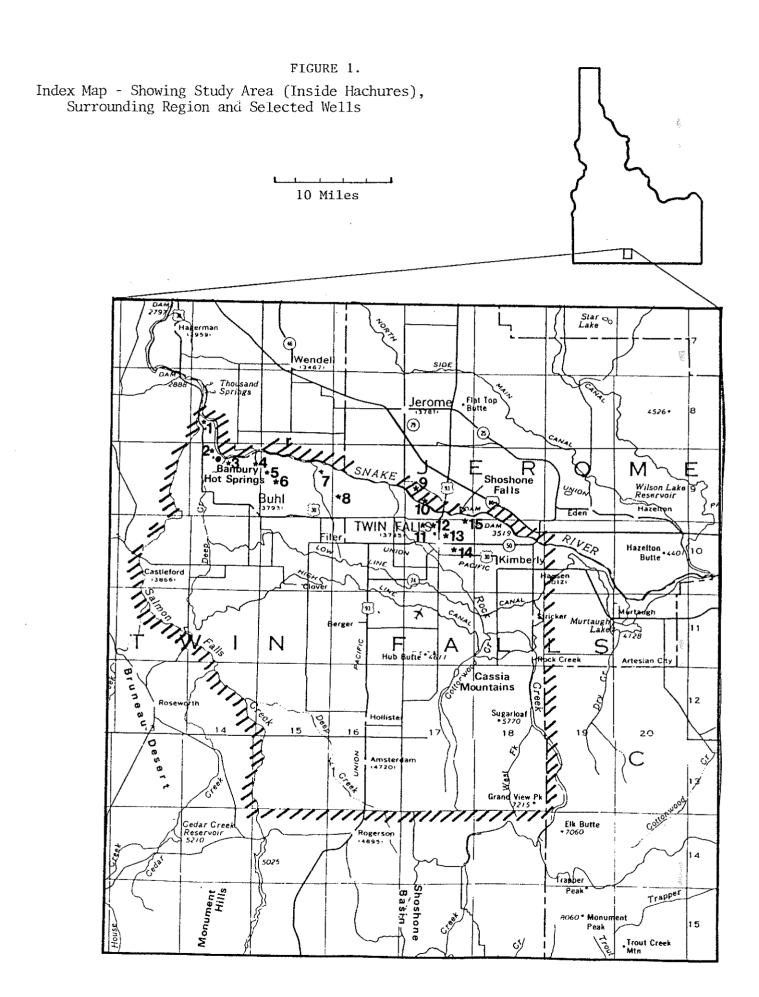
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INTRODUCTION

The low to moderate geothermal resource base of the Snake River Plain, Idaho is extensive. Estimates of mean reservoir energy of over 455 x 10¹⁸ J have been made by Mabey (1983) and energy offset of over 26 MW thermal were made by Bloomquist and others (1985). The hydrothermal resources range in temperature from 30 to 70° C (86 to 158°F) and well depths range between 80 and 3000 feet. Most flowing wells have high yields (2000 gpm) and shut-in pressures (250 psi). The focus of this report is the geothermal resources of southern Idaho's Twin Falls County, Figure 1.

Artifacts found in the vicinity of Banbury and other hot springs in Twin Falls County are evidence that thermal waters were known and used by Indians and early settlers. In 1938, a 1154 foot well was drilled by the City of Twin Falls (T10S, R17E, SESWSW) that encountered the geothermal resource. well continues to be used for irrigation. Further development did not ensue until the early 1970's when concern for energy was Drilling continued until the early 1980's when proheightened. jects for low-head hydro-power production, space heating, irrigation and aquaculture were developed. The increased demands on the resource have resulted in noticeable declines in pressure in existing artesian wells. These declines have caused concern among the resource owners, in some cases have and litigation.

In the study area, there are more than seventy-five thermal wells. For the convenience of discussion and ease of reference, the wells have been assigned an arbitrary number. The locations of selected wells referred to in this study are shown on Figure 1 and driller's logs for these wells and others are found in Appendix A.



Purpose and Approach

This study was initiated to provide baseline resource data for the thermal system in Twin Falls County and to develop a model that would describe the relationship between the geothermal resource and the regional geologic framework. Methods selected to complete the objectives were: 1) compile data for existing wells; 2) analyze well logs for thermal water levels and subsurface geology; 3) compile and interpret geologic mapping in relation to regional stratigraphy and structural controls of the system and possible recharge areas; 4) establish a monitoring network using existing thermal wells to record pressure and temperature measurements on a regular basis; and 5) collect samples of the suspected rhyolite reservoir rock for geochemical analysis and comparison.

Previous Work

Geologic Studies

Geologic mapping within the study area has been limited to reconnaissance studies. For the southern part of Twin Falls County, virtually no published geologic maps exist at a scale of less than 1:250,000. Also, the nature and stratigraphy of the Idavada rhyolites exposed throughout the area are largely unknown.

The early geologic work in the Snake River Plain included reconnaissance level studies of geology by King (1878), Lindgren (1898), Russell (1902), Schrader (1912), and Buwalda (1924). Stearns and others (1938) published a report on eastern Snake River Plain geology and groundwater resources that included descriptions of the basalts of the lava plains and the Tertiary volcanic rocks at Shoshone Falls. Stearns, in a later publication (1955), described the mud flow associated with the Shoshone Falls volcanic rocks. Youngquist and Haegele (1956) published a

report and very general geologic map of the Cassia Mountains. Malde and Powers (1962) studied and defined the stratigraphic framework for the western Snake River Plain and developed much of the stratigraphic terminology and nomenclature currently in use. As part of their reconnaissance of the western Snake River Plain, Malde and others (1963) included much of the western portion of the study area on a geologic map at a scale of 1:125,000. and Powers (1972) published a more detailed geologic map of a part of the previously mapped area, which included the Banbury Hot Springs area, at a scale of 1:48,000. Covington (1976), published a 1:24,000 scale geologic map which focused on geologic hazards along the Snake River Canyon in the Twin Falls area. Stratigraphic relations of Paleozoic and Mesozoic marine sedimentary rocks in the Cassia Mountains were presented by Mytton and others (1983). A study and morphometric analysis of basalt volin the southern and western portion of the study area was canoes conducted by Jenks in 1984.

Geothermal Studies

The geothermal resources of the Banbury and Artesian City areas were first described by Stearns, and others (1938). Young and Mitchell (1973) sampled hot springs and wells in Twin Falls County as part of a statewide reconnaissance of geologic setting and chemical characteristics of thermal waters. Lewis and Young in 1980 and 1987 (in press) investigated hydrologic conditions of thermal waters in the Banbury and Twin Falls area using geophysical studies and water chemistry. The geothermal resources in the Artesian City area were studied by Struhsacker and others (1983).

Climate

The climate of the area is characterized by hot, dry summers [>32°C(90°F)] and cold winters [<7°C(20°F)]. The mean annual temperature at Twin Falls is 9°C(49°F) (U.S. Department of Commerce, 1973). Mean annual precipitation ranges from 9.5 inches at Twin Falls (U.S. Department of Commerce, 1973) to over 40 inches (Wilson and Carstens, 1975) in the mountainous areas to the south. The majority of the precipitation falls as rain or snow during the winter.

Physiography

Twin Falls - Banbury Study Area is within the Snake The River Plain subdivision of the Columbia Plateaus physiographic province (Fenneman, 1946). The Snake River Plain is an arcuate structural and topographic depression extending across southern Idaho from Weiser on the west to Yellowstone National Park on the Recent workers have further subdivided the plain as consisting of a western graben, eastern downwarp, and central transition zone (Mabey, 1982). The Twin Falls - Banbury Study Area lies within the transition zone between the eastern and western Topography of the plain is subdued and includes subdivisions. extensive lava plains and numerous constructional volcanic including basaltic cinder cones and broad shield features Elevations of the Snake River Plain range from 2,000 volcanoes. feet on the west to 5,000 feet near the eastern end.

The area in and around Twin Falls County, like much of the eastern Snake River Plain, is characterized by a gently rolling lava plain from which rise scattered shield volcanoes such as Hansen, Stricker, Hub, Sonnichsen and Flat Top Buttes (see Plate 1). These volcanic shields generally rise from 150 to 400 feet above the surrounding plain (Jenks, 1984). Many of the younger Pleistocene lavas on the north side of the Snake River Canyon

still preserve features such as pressure ridges and aa and pahoehoe lava surfaces. On the south side of the river the lavas are earlier Pleistocene and Pliocene and most original features have been removed by erosion or obscured by loess (Covington, 1976).

Within the study area the plain extends south from the Snake River approximately 15 miles to the Cassia Mountains (locally known as the South Hills or Rock Creek Hills). The Cassia Mountains are a broad domal uplift that rise from a base elevation of 4200 feet at the point where Rock Creek enters the plain to a maximum of over 8000 feet at Monument Peak. The uplift contains a core of Paleozoic marine sedimentary rocks that are unconformably mantled by a sequence of outwardly dipping Tertiary ash-flow tuffs. The dome is deeply dissected by numerous drainages.

The southwest and western portion of the Twin Falls - Banbury Study Area is a gently rolling lava plain broken by low shield volcanoes. In addition, these areas are transected by a series of low (50 to 100 foot) northwest trending fault scarps and low elongate shield-like ridges formed by basaltic fissure eruptions.

The Snake River is the primary drainage for the study area. Through the Twin Falls reach, the river has incised a nearly vertical walled canyon 500 to 600 feet deep and 1475 to 4900 feet in width. Important tributaries within the study area include Rock Creek and Salmon Falls Creek. Rock Creek is the primary drainage of the Cassia Mountains where it has carved a canyon in excess of 1450 feet deep. Numerous other less dramatic but deeply incised drainages issue from the Cassia Mountains in a radial pattern. Between the Cassia Mountains and the Snake Rock Creek Canyon is relatively shallow (30 to 50 feet) River, but steep walled. The canyon of Salmon Falls Creek, which borders the study area on the west, is deeply incised with depths in excess of 500 feet. The remaining drainage for the area consists of a series of shallow drainageways that have developed along fault scarps and in the creases between lava flows and shield volcanoes.

GEOLOGIC FRAMEWORK

Rocks exposed in the Snake River Plain are predominantly the products of bimodal basalt-rhyolite volcanism during the past 12 million years (Leeman, 1982). These volcanics were erupted onto a region of probable low relief which was underlain by folded and faulted pre-Cenozoic sedimentary, volcanic, and plutonic rocks (Armstrong, 1978). Inception of volcanism in any one area of the plain was characterized by voluminous explosive eruptions of rhyolitic ash which resulted in extensive sheets of welded tuff. were followed by extrusions of rhyolite lavas with some minor intercalated basalts. The final phase of volcanism for an area was the eruption of basalt lavas. Based on radiometric dating (Armstrong, 1975; Armstrong and others, 1975; Armstrong and others, 1980; Honjo and others, 1986), volcanism has progressed along the axis of the eastern plain from the Owyhee Plateau area on the southwest to Yellowstone on the northeast. including measurements of flow directions in Recent evidence, welded tuffs along the margin of the plain (DeTar and Street, unpub. data; Leeman, unpub. data) and additional radiometric dating (Honjo and others, 1986), points to the probability that this general progression was interrupted by regressions or coeval activity at more than one eruptive center.

Stratigraphy

General Statement

The oldest rocks in the study area are Paleozoic and Mesozoic marine sediments that form the core of the Cassia Mountains southern margin of the Snake River Plain (Rember and Bennett, 1979; Mytton and others, 1983). Tertiary rhyolitic volrocks of the Idavada group mantle the pre-Cenozoic sediments in the Cassia Mountains and underlie most of the Snake Tertiary rhyolites comprise the known basement basalts. rock within the Snake River Plain portion of the study area and are the reservoir rocks for the Twin Falls - Banbury geothermal Tertiary rhyolite is overlain by 500-700 feet of The Tertiary and Quaternary basalts. Lacustrine and fluvial sedimentary rocks of the Glenns Ferry and Bruneau formations are interbedded with and overlie the basalts in the western portion of the study area (Malde and Powers, 1962). The following description the major rock units focuses on rocks of the Idavada Volcanic Group because of their relationship to the Twin Falls - Banbury geothermal system. Descriptions of the other rock units are more generalized and are included to provide a more complete picture of the stratigraphic framework of the study area.

Pre-Cenozoic Rocks (MPz)

Paleozoic and Mesozoic sedimentary rocks are exposed both on the northern and southern margins of the eastern Snake River Plain and probably formed a continuous terrain across the area prior to the inception of the Miocene rhyolite volcanism (Armstrong, 1978). Within the study area, the pre-Cenozoic sediments are exposed in the core of the Cassia Mountain uplift. Based on these exposures the pre-volcanic topography was probably

a gently rolling to mountainous upland (Struhsacker and others, 1983). Numerous pre-Tertiary episodes of tectonic deformation have resulted in the sediments being folded and thrust faulted.

The units in the Cassia Mountains consist of limestone, dolomitic limestone, siltstone, quartzite, and chert. The limestones and dolomitic limestones range from light to dark gray and fine to medium grained and occur in thick massive beds several to several tens of feet thick. Chert occurs as thin interbeds, as major members of formations, and as nodules and pods up to several feet across within limestone units. Siltstones show a wide range of colors including reddish brown, yellow, and gray, and commonly occur as a series of thin laminar beds. Quartzites are buff, reddish and gray-green, fine to medium grained and massively bedded.

Tertiary Idavada Volcanic Group

The oldest volcanic rocks in the study area have been mapped part of the Tertiary Idavada Volcanic Group (Rember and Bennett, 1979; Malde and Powers, 1962). The Idavada Group is a loosely defined aggregation of predominately rhyolitic welded ash-flow tuffs with subordinate amounts of rhyolite lava and intercalated tuffaceous lacustrine sediments (Malde and Powers, Rocks of the Idavada Group are exposed in the study area within the Snake River Canyon near the City of Twin Falls, in the Cassia Mountains 10 to 15 miles to the south, and in various stretches of Salmon Falls Creek to the west. Idavada rocks have also been mapped in areas along the northern margin of the Snake River Plain (Malde and others, 1963). While rocks of the Idavada group are clearly related genetically, they probably represent numerous as yet undifferentiated episodes of volcanism ranging in age from 12 million years to as young as 6 million years (B. Bonnichsen, personal communication).

Idavada Ash-Flow Tuffs in the Cassia Mountains (Tiv)

Unconformably overlying the Paleozoic and Mesozoic sediments in the Cassia Mountains are a sequence of welded vitric and vitric-crystal rhyolite ash-flow tuff sheets of Miocene age (J. Mytton and others, unpub. geologic map). The sequence consists of seven or more distinct cooling units, most of which are separated by variable amounts of poorly consolidated airfall, water-lain, and non-welded ash-flow tuff (DeTar and Street, unpublished data). Thicknesses of the welded tuffs range from 15 feet to over 200 feet. The densely welded units are distinctly zoned, consisting, in ascending order, of: (1) gray bedded surge deposits; (2) gray to black basal vitrophyre; (3) brown, reddishbrown, red and purple massive to platy thick densely welded central lithoidal zones; and (4) an upper lithoidal zone containing abundant primary flow marks and secondary flow folds. Thin gray to black vitrophyre also occurs irregularly at the tops of some Phenocryst assemblages and percentages are variable from unit to unit but are generally dominated by plagioclase with amounts of pyroxene, quartz, and opaque oxides (probably magnetite and ilmenite) (DeTar and Street, unpub. data).

A generalized stratigraphic section for the Cassia Mountains sequence is presented as Appendix B and is based on a reconnaissance level investigation of exposures along Goat Springs Creek (T13S, R17E, Sec 17) and Dry Gulch (T12S, R18E, Sec 2 & 3).

Older Rhyolite (Not Exposed)

All of the wells in the Snake River Plain portion of the study area that have been drilled deep enough to completely penetrate the Cenozoic basalts encounter Idavada rhyolite. Data for these rhyolites are generally meager, consisting primarily of well cuttings, driller's logs, and discussions with drillers. The descriptions in this report are based primarily on information

from wells. Direct correlation of these units with the welded tuff sequence of the Cassia Mountains is not possible because of the extensive basalt cover between the wells and the mountains, the possibility of intervening subsurface structures between the two areas, and the lack of detailed stratigraphic information for the subsurface units in the study area.

Cuttings from the USGS test well at Filer were examined for this study. The cuttings were sand to pebble size making it impossible to determine the large scale features of the rocks. However, it was possible to identify vitric and lithoidal phases rhyolitic composition. Constituents include abundant glass shards and clear glassy feldspar. Glass shards are cuspate and platy and commonly show bubble wall structure. Other constitinclude black and brown 0.01 inch rounded obsidian fragments which are similar to the "apache tears" that commonly occur in perlitic vitrophyres of the Idavada welded tuffs in the Cassia Mountains. Fragments in some intervals, both crystals and glass, are partially to completely coated with a white chalky encrusta-Also noted in the cuttings were minor amounts (less than 1%) of pumice fragments and masses of glass shards welded together.

Lacustrine Sediments

Wells drilled in the Filer (No. 8, Figure 1) and the City of Twin Falls area (No.'s 11, 12, 13, 14 and 15) encounter a 50 to 100 foot section of white to tan tuffaceous probable lacustrine siltstone and clay-stone overlying the Idavada rhyolites. Cuttings from the USGS test well contain abundant calcareous oolites or coated grain sands similar to exposed Tertiary and Quaternary lacustrine sediments in the western Snake River Plain (Swirydczuk and others, 1979; W. Burnham, personal communication). The oolites or coated sand grains make up approximately 20 to 40 percent of the sample.

The lower part of the lacustrine section as encountered in the USGS test well is a white sugary, vitreous, highly silicified unit containing scattered 0.02 to 0.03 inch masses of opaque oxides, some of which have a halo of limonite. Because the samples were limited in quality and quantity, little else is known about this unit.

Subsurface Alteration and Mineralization

rhyolite and lacustrine sediments Ιn the subsurface described above. evidence of alteration and mineralization was noted including the white chalky encrustation, fragments and coatings of opal and chalcedony occurring within rhyolite, and magnetite within a probable silicified sedimentary unit. opal, chalcedony, magnetite, and silicification are the products of hydrothermal activity. However, based on the existing samples it is not possible to determine whether the occurrences are related to the present thermal system or to hydrothermal activity associated with volcanism. The white chalky encrustation may represent paleo-episodes of duricrust formation.

Shoshone Falls Rhyolite

A silicic volcanic unit of Idavada affinity is exposed in the Snake River Canyon from the Canyon Springs Golf Course (T9S, R17E, Sec 33 NW) to approximately 1/2 mile west of Twin Falls dam (T10S, R18E, Sec 4) and is identified as a single rhyolite lava flow. The rock is light to dark gray on fresh surfaces and weathers to a dark reddish brown. Fracturing occurs throughout the unit. Zones of sheeted or platy fractures are common. Sheets or plates within the zones are generally from 0.5 to 2 inches thick and the zones may be quite extensive covering

several hundreds of square feet. Columnar jointing isn't pronounced, but strong vertical fractures are abundant. The bottom of the flow is not exposed but well drilling has shown the flow to be as much as 600 feet thick.

Microscopic examination, as reported by Stearns and others, (1938), and confirmed in this study, has revealed no definitive criteria on which to distinguish this unit as a lava or tuff. Bonnichsen (1982 a & b) has found that many of the tuffs and of the Idavada group are petrographically indistinguishable. Phenocrysts in the porphyritic rock are predominately plagioclase with minor amounts of pyroxene and opaque oxides. plagioclases are milky white in hand specimen and are up to 0.2 inch in length. The pyroxenes and oxides are commonly closely associated and phenocryst aggregates of plagioclase, pyroxene and oxides, are common but widely scattered. Bonnichsen (1982b), notes that phenocryst aggregates in general tend to be more abundant in lava flow rocks than in welded tuffs, but more detailed statistical analysis will be needed to establish the greater abundance of the aggregates in this unit.

These rocks have been classified (Stearns, 1955; Stearns and others, 1938) as an andesite lava flow. Classification of this unit as a lava flow was based on the abundance of glass near the top of the unit, the abundant vesicles, and the common flow structures. However, all of these characteristics can also be found in the ash-flow tuffs of the Idavada group in areas north and south of the Snake River Plain. Further evidence suggesting that this unit is a lava flow rather than an ash-flow tuff, is the lobate morphology, irregular upper surface of the unit, the irregular distribution of vitrophyre and breccia zones within the unit, the lack of typical ash-flow tuff zonation, and the lack of the characteristic ash-flow tuff tabular aerial distribution.

While Stearns' classification of this unit as a lava flow can be supported, it is believed his classification of the rock as andesite is incorrect within the context of current classification schemes. Based on the phenocryst assemblage, which is typical of Idavada group rocks, and geochemical analyses (see Appendix C, samples SF1 and SF2) these rocks should be classified as rhyolite or silicic latite genetically related to the Idavada group.

Tertiary and Quaternary Basalts (Tb, QTb and Qb)

A sequence of Tertiary and Quaternary basalt lava flows overlies the Tertiary rhyolites. The oldest of these basalts type Banbury Basalt of Stearns and others, (1938) includes the which has yielded radiometric ages of 4.4 to 4.9 million years and possibly as old as 6.2 million years, (Armstrong, 1975, contains a good discussion of Banbury Basalt and its various ages and correlations). Near the type locality (Stearns and others, Banbury Basalt lies unconformably on Idavada rhyolitic rocks (Malde and Powers, 1962). While contact between Banbury and Idavada rocks is not exposed in the Snake River Canyon near the City of Twin Falls, Banbury Basalt is believed to overlie Idavada volcanics over much of the area to the west. where the Shoshone Falls rhyolite is exposed in the Snake River Canyon it is not overlain by Banbury Basalt but rather by younger basalts of Glenns Ferry (Plio-Pleistocene) and Snake River (Pleistocene-Holocene) age (Covington, 1976). Therefore, the rhyolite marks the eastern limit of Tertiary Shoshone Falls basalt flows in this area. Banbury Basalt is also exposed in the uplands of the Snake River Plain in the southwest part of the study area and within the Salmon Falls Creek Canyon (Malde and others, 1963).

Banbury Basalt within the study area consists of a series of thin olivine basalt flows which are locally interbedded with minor stream and lake sediments (Malde and Powers, 1972). The basalts are gray to dark gray to gray-green, both dense and vesicular, fine to coarse grained and diktytaxitic (Malde and others, 1963). These older basalts generally are altered. Phenocrysts include amber to brown olivine crystals and clusters of plagioclase 0.01 inch in diameter and laths of plagioclase averaging 0.08 inch in length.

A series of similar appearing, but younger and unaltered basalts, overlie Banbury Basalt in the City of Twin Falls area. These basalts include units assigned to the Glenns Ferry Formation (Malde and others, 1962 and 1963) in the western part of the study area and the younger Snake River Group (Malde and others, 1962; Covington, 1976) of lava flows which were erupted from vents north, south and northeast of Twin Falls. These basalts represent the surface units at the City of Twin Falls and are well exposed along the canyon wall of the Snake River and Rock Creek.

Geochemistry of Idavada Rhyolites

In order to establish baseline geochemistry for the known and related geothermal reservoir rocks of the area, nineteen samples of ash-flow tuff, airfall tuff and rhyolite lava were analyzed for major element compositions. Of these, nine samples were analyzed for the following trace elements: Sr, Ba, Co, Cu Pb, Zn, Sb, Li, Be, Fr, La, Ce, and F. The variable compositions of the rocks fall within a range comparable to other rhyolitic rocks of the Idavada Volcanic Group (Bonnichsen, 1982a and b; Bonnichsen and Citron, 1982; Ekren and others, 1982; Leeman, 1982). A discussion of analytical techniques and the results of the analyses are presented in Appendix C.

Tectonics and Structure

Mabey (1982) states that "...the Snake River Plain remains one of the least understood geologic structures in the United It has been described as a depression, downwarp, graben, rift, and lateral rift." The plain is clearly related to, and the partial result of the propagation of, bimodal basalt-rhyolite volcanism from southwest to northeast across southern Idaho. Mabey (1982) divides the plain into three subdivisions; an eastern, western, and central segment. The western Snake River Plain graben is clearly expressed topographically and is distinctly bounded by normal faults along its southwest and northeast margins. The eastern Snake River Plain, Mabey concludes, is "likely a downwarp containing a complex of calderas." Ekren and others (1984) considered the eastern Snake River Plain to be part of a tectonic-structural trend extending from the Idaho-Oregon-Nevada border area to Yellowstone National Park and possibly beyond. Within the central segment, graben bounding structures of the western Snake River Plain, but not the pronounced topographic expression, are superimposed on the eastern Snake River Plain trend. The Twin Falls - Banbury Study Area is located within this central segment.

Not only is the gross structure of the Snake River Plain poorly understood, but the local structures within the plain are similarly enigmatic. Basement and subsurface structures are probably complex because of the complex geologic and volcanic history of the area. However, these structures are difficult to completely delineate because they are covered by Cenozoic basalts.

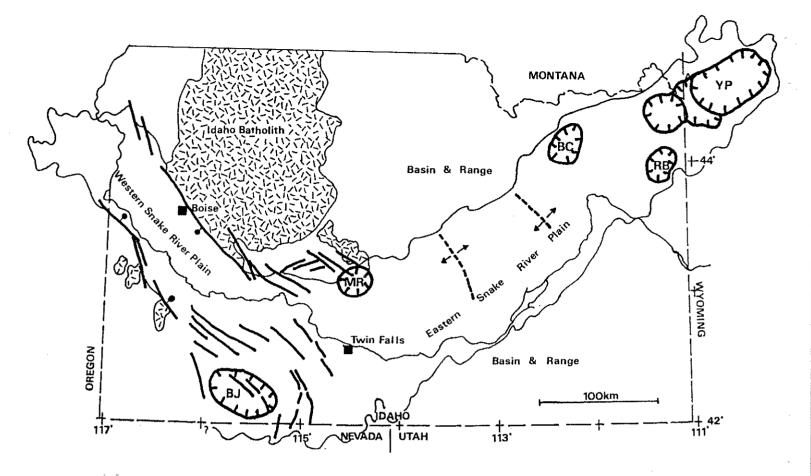
Structural features characteristic of both the eastern and western Snake River Plain occur within the study area. Idavada volcanic strata along the southern margin of the plain, within the study area, dip into the plain and disappear beneath basalt

without any clear evidence of down-faulting. A review of the depth at which the top of Idavada rocks is encountered in wells south of the City of Twin Falls demonstrates that this trend continues beneath the basalt cover (Street and DeTar, unpub. data). This is consistent with interpretation of the eastern Snake River Plain as a downwarp.

Western Snake River Plain structures are represented within southwest part of the study area by a series of northwest trending small displacement (<100 ft.) normal faults and small scale grabens that fan out from the Salmon Falls Creek area across the Bruneau Desert to merge with the fault zone on the southwest margin of the western plain at the base of the Owyhee Mountains (see Figure 2 and Plate 1). These northwest-trending fault zones are characteristic of the structural style within the western part of the study area. The northwest trending faults of the Bruneau Desert and Salmon Falls Creek area form a broad zone of normal faulting which appears to be related to Basin and Range extension. This zone crosses the southwest portion of the study area from Hollister to Buhl and is sub-parallel to rift zones normal to the eastern plain. The rift zones are also thought to be related to Basin and Range extension, (Kuntz, 1977).

The northeastern most trend of the Bruneau Desert - Salmon Falls Creek zone is here informally designated the Buhl-Berger Structure Zone (BBSZ). This N 35°W trending structure zone traverses the area from southeast of Buhl to southeast of Hollister, and is delineated based on the occurrence and trend of fault scarps and associated basaltic eruption fissures, and aligned vents (see Plate 1).

The principal fault scarps within the BBSZ are recognized in the field and on topographic maps as somewhat rounded northwest trending linear ridges which are steepest on the northeast side and slope gradually away to the west indicating a northeast side down sense of movement. There are two primary scarps within the



Compiled from Ekren and Others (1982), Protska and Embree (1978) and Leeman (1982).



Known or Inferred Calderas or Eruptive Centers

Caldera Designations

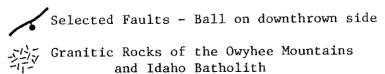
YP - Yellowstone Complex

RB - Rexburg Complex

MR - Magic Reservoir Caldera

BJ - Bruneau-Jarbidge Eruptive Center

BC - Blue Creek Caldera



Rift Zone

Figure 2
Tectonics and Structure

zone, one, 50 to 75 feet high, extending from the southeast corner of Buhl to about Berger (Plate 1) and the second, 150 to 175 feet high, from Clover to just north of Hollister (Malde and others, 1963).

Basaltic fissure eruptions are associated with both of the fault scarp segments, evidence that the BBSZ is probably a deep seated structure. Southeast of Hollister, three aligned basaltic shield volcanoes lie on trend with the BBSZ and appear to represent an extension of the zone. The zone cannot be traced with any certainty into the Cassia Mountains. A graben which forms the Shoshone Basin (Figure 1) is a possible southeast extension of the zone, but this interpretation is complicated by a possible east-northeast trending structural zone that has been identified trending across the northern margin of the Monument Hills and into a zone of concentrated and possibly faulted basaltic vents including Salmon Falls Butte (Figure 1), (Covington, Geological Survey, unpublished mapping; Bonnichsen & Jenks, Idaho Geological Survey, personal communications). This east-northeast may offset structures in the Cassia Mountains and trending zone Monument Hills (Figure 1) from those in the plain and it appears to at least mark a change in structural style and trend.

On the northwest the BBSZ can be traced to a zone of north-west trending faults mapped by Malde and others, (1963) through the Banbury Hot Springs - Melon Valley area. The BBSZ and the extension mapped by Malde and others, (1963) coincide with a major bend in the Snake River and the sense of movement (down to the northeast) is appropriate for fault control of this deflection of the river.

In some areas of the Snake River Plain alignment of basaltic cinder cones and/or shields can be used to delineate structures such as the BBSZ. Probably most basalt volcanoes in the Snake

River Plain are related to deep seated fractures. However, basaltic vents in the area northeast of the BBSZ area show no clear pattern of alignment.

The stratigraphy of the study area is summarized in Figure 3. Structurally, the area is complex with a dominant northwest-trend. The stratigraphy and structures both play important roles in understanding the nature and occurrence of the geothermal resource. The following hydrogeology sections discuss the relationship between the geothermal resource and the geologic framework.

HYDROLOGY OF THERMAL SYSTEMS

Aquifer Characteristics

The geothermal aquifer of the Twin Falls - Banbury area is a confined artesian system with shut-in pressures ranging from 14 to 250 psi depending upon the elevation of the well. Based upon the drillers' logs, well cuttings and chemistry of thermal water, it is apparent that the system is contained in the Idavada volcanics with the upper units acting as the confining layer. Permeability of the volcanic rocks is the result of structural fractures related to tectonic movements, sheeted cooling joints and fractures developed during emplacement, intergranular porosity of the non-welded ash flows and air fall tuffs, and voids left between successive flows.

Where Idavada pyroclastic rocks are exposed, they are composed of a heterogeneous assemblage of rocks. In heterogeneous rocks, it is difficult to accurately calculate transmissivity and storativity values without long-term flow tests. The following section describes the data obtained from the only long-term flow test on the aquifer near Twin Falls.

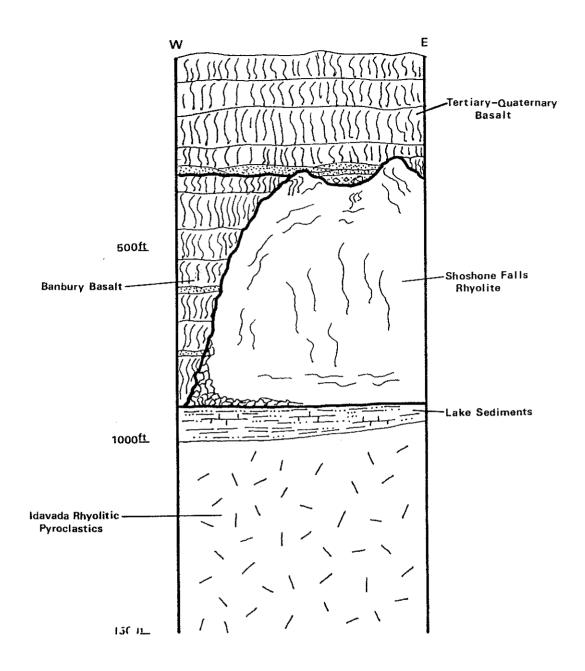


Figure 3
Generalized Stratigraphy of the Study Area

Aquifer Testing

A 1000 hour (41.7 days) aquifer test was conducted by CH₂M Hill (unpub. report, 1982) for the developer of the well in T9S, R17E, Sec 33 NWNWNW (Well No. 10) from July 22 to September 2, 1982. During the test, the well-head pressure was maintained at 150 psi which was the pressure that was to be used for production. Periodic measurements of the flow indicated that it remained relatively constant at 2,920 GPM throughout the test.

During the test, shut-in pressures were recorded on the wells in T10S, R17E, Sec 4 (Wells No.'s 11 and 12). At the end of the test, pressures were recorded in the two monitoring wells and the test well for 46 days to determine the rate of recovery. Nearly complete recovery to original static heads was noted in all three wells.

Based on this test, the transmissivity (permeability) and storativity (volume of water taken into or released from storage) were calculated for both recovery and drawdown and are as follows:

Well I	Data Type	nsmissivity (GPD/FT) Storativit	<u>Storativity</u>		
9S 17E 33 #10	Recovery	74,300 -			
10S 17E 4 #12	Recovery	46,400 -	,		
	Drawdown	45,200 5.8 x 10 ⁻⁴	±		
10S 17E 4 #11	Recovery	58,600 -	А		
	Drawdown	$44,600$ 6.2 x 10^{-4}	+		

From this test and subsequent time drawdown analyses, $\mathrm{CH}_2\mathrm{M}$ Hill concluded that no hydrologic boundaries exist between the wells and that the development of the well would affect the water levels in nearby thermal wells.

Water Chemistry

Detailed analyses of thermal water chemistry were not conducted during this study. Previous analyses have shown that the thermal water in the study area is low in Ca, Mg and high in Si, HCO₃ and F in comparison to other geothermal waters in the state with different aquifer rock types (Young and Mitchell, 1973).

It has been determined that the Idavada ash flow tuffs are high in SiO₂, F and low in Ca and Mg (Appendix C). It is apparent from this that the chemistry of the thermal water reflects the geochemistry of the flowpath and Idavada host rock.

Monitoring of Thermal Wells

A monitoring network of five wells in the Banbury Hot Springs area was established in the fall of 1983. A similar network of four wells was established for the Twin Falls area in the spring of 1984. A standard method was employed to determine pressure and temperature measurements. The procedure included completely shutting in each well for ten minutes and at the end of this time recording pressure and temperatures. Record was also kept of other information, such as how many valves were opened or comments made by the owner regarding well performance.

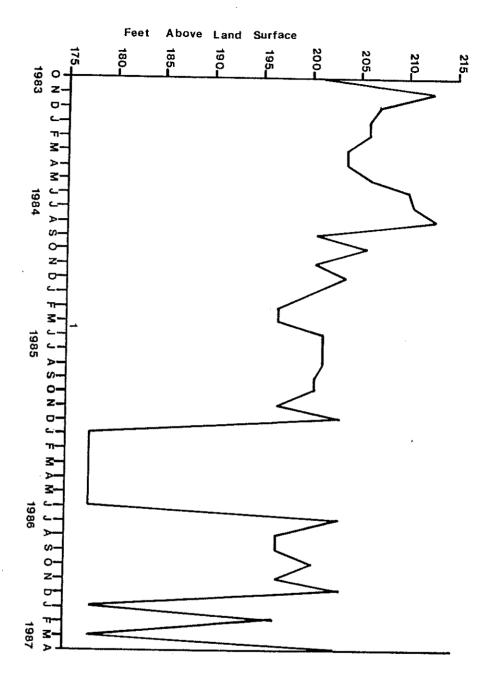
The locations of the monitoring wells and other wells discussed in this section are shown in Figure 1. Frequency of measurements and pertinent well information for both the monitored wells and other wells discussed in the text are found in Table 1. The results of the monitoring effort are illustrated by hydrographs (Figures 4 through 12). Temperatures remained constant during the monitoring.

TABLE 1. WELL INFORMATION AND FREQUENCY OF MEASUREMENTS

FLUID

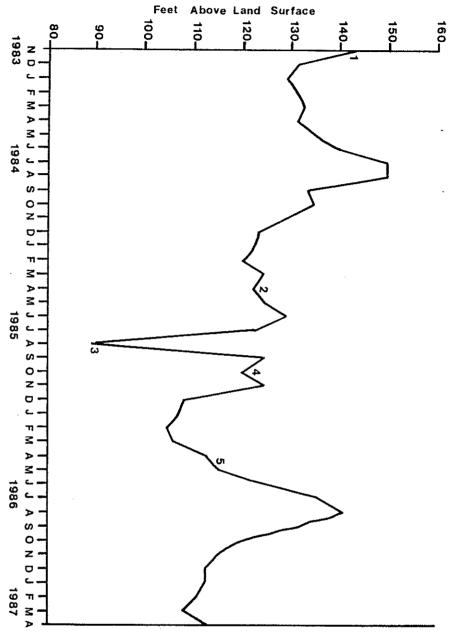
				ELEVA-	TEMP.		*FRE-	
	LOCAT	rion		TION	°C/°F	DEPTH	QUENCY	PRESENT USE
88	14E 3	30 NES	SWSE	2897	68°/154°	760 <i>′</i>	M	Heating
88	14E 3	33 SW1	NWSE	2907	47°/117°	800′	М	Greenhouse Heating
9S	14E (04 NWI	WNW	2959	41°/106°	700 <i>′</i>	М	Space Heating
9S	14E (09 NES	SESW	2996	32°/90°	850′	M	Tropical Fish
9S	14E	14 NWS	SENW	3146	34°/93°	906′	M	Irrigation
9s	14E 3	14 NEI	NESW	3200		1300′		Catfish Propogation
9S	15E :	12 SW:	SWNE	3064	42°/108°	1420′	М	Irrigation
9s	16E 2	20 NE:	SESE	3530		1247′		USGS-Observation Well
9S	17E 2	29 NW:	SWSE	3154		743′		Cat Fish Propogation
9S	17E 3	33 NWI	WNW	3174		750 <i>′</i>		Low-Head Power Gen.
10S	17E (04 SWI	NWNE	3662	37°/99°	1453′	W	Campus Space Heating
10S	17E (04 SW	SENE	3668	37°/99°	1191′	W	Campus Space Heating
10S	17E :	10 NW:	SENW	3717	37°/99°	1700′	V	Space Heating
10S	17E	14 SW	SESE	3786		1154′		Irrigation
10S	18E (o6 nwi	WNW	3585		1300′		Domestic Space Heating
	85 95 95 95 95 95 95 105 105 105	8S 14E : 8S 14E : 9S 14E : 9S 14E : 9S 14E : 9S 15E : 9S 16E : 9S 17E : 10S 17E : 10S 17E : 10S 17E :	8S 14E 33 SWI 9S 14E 04 NWI 9S 14E 09 NES 9S 14E 14 NWI 9S 14E 14 NES 9S 15E 12 SWI 9S 16E 20 NES 9S 17E 29 NWI 9S 17E 33 NWI 10S 17E 04 SWI 10S 17E 04 SWI 10S 17E 10 NWI 10S 17E 14 SWI	8S 14E 30 NESWSE 8S 14E 33 SWNWSE 9S 14E 04 NWNWNW 9S 14E 09 NESESW 9S 14E 14 NWSENW 9S 15E 12 SWSWNE 9S 15E 12 SWSWNE 9S 16E 20 NESESE 9S 17E 29 NWSWSE 9S 17E 33 NWNWNW 10S 17E 04 SWNWNE 10S 17E 04 SWSENE 10S 17E 04 SWSENE	LOCATION TION . 8S 14E 30 NESWSE 2897 8S 14E 33 SWNWSE 2907 9S 14E 04 NWNWNW 2959 9S 14E 09 NESESW 2996 9S 14E 14 NWSENW 3146 9S 14E 14 NENESW 3200 9S 15E 12 SWSWNE 3064 9S 16E 20 NESESE 3530 9S 17E 29 NWSWSE 3154 9S 17E 33 NWNWNW 3174 10S 17E 04 SWNWNE 3662 10S 17E 04 SWSENE 3668 10S 17E 10 NWSENW 3717 10S 17E 14 SWSESE 3786	LOCATION TION °C/°F 8S 14E 30 NESWSE 2897 68°/154° 8S 14E 33 SWNWSE 2907 47°/117° 9S 14E 04 NWNWNW 2959 41°/106° 9S 14E 09 NESESW 2996 32°/90° 9S 14E 14 NWSENW 3146 34°/93° 9S 14E 14 NENESW 3200 9S 15E 12 SWSWNE 3064 42°/108° 9S 16E 20 NESESE 3530 9S 17E 29 NWSWSE 3154 9S 17E 33 NWNWNW 3174 10S 17E 04 SWNWNE 3662 37°/99° 10S 17E 04 SWSENE 3668 37°/99° 10S 17E 10 NWSENW 3717 37°/99° 10S 17E 14 SWSESE 3786	BS 14E 30 NESWSE 2897 68°/154° 760′ 8S 14E 33 SWNWSE 2907 47°/117° 800′ 9S 14E 04 NWNWNW 2959 41°/106° 700′ 9S 14E 09 NESESW 2996 32°/90° 850′ 9S 14E 14 NWSENW 3146 34°/93° 906′ 9S 14E 14 NENESW 3200 1300′ 9S 15E 12 SWSWNE 3064 42°/108° 1420′ 9S 16E 20 NESESE 3530 1247′ 9S 17E 29 NWSWSE 3154 743′ 9S 17E 33 NWNWNW 3174 750′ 10S 17E 04 SWSENE 3662 37°/99° 1453′ 10S 17E 04 SWSENE 3668 37°/99° 1191′ 10S 17E 10 NWSENW 3717 37°/99° 1700′ 10S 17E 14 SWSESE 3786	LOCATION TION °C/°F DEPTH QUENCY 8S 14E 30 NESWSE 2897 68°/154° 760′ M 8S 14E 33 SWNWSE 2907 47°/117° 800′ M 9S 14E 04 NWNWNW 2959 41°/106° 700′ M 9S 14E 09 NESESW 2996 32°/90° 850′ M 9S 14E 14 NWSENW 3146 34°/93° 906′ M 9S 14E 14 NENESW 3200 1300′ 9S 15E 12 SWSWNE 3064 42°/108° 1420′ M 9S 16E 20 NESESE 3530 1247′ 9S 17E 29 NWSWSE 3154 743′ 9S 17E 29 NWSWSE 3154 743′ 9S 17E 04 SWNWNW 3174 750′ 10S 17E 04 SWSENE 3668 37°/99° 1453′ W 10S 17E 10 NWSENW 3717 37°/99° 1700′ W 10S 17E 10 NWSENW 3717 37°/99° 1700′ W

^{*} Frequency of Measurements (M=monthly, W+weekly)



1. No reading possible for April and May

Figure 4
Well #1 Hydrograph
T8S, R14E, Sec. 30, NESWSE



- Well not in use 1.
- Greenhouses on line 2.
- 3.
- 4.
- New well being drilled Developing system adding greenhouses Used for domestic from April through July

Figure 5

Well #2 Hydrograph

T8S, R14E, Sec. 33, SWNWSE

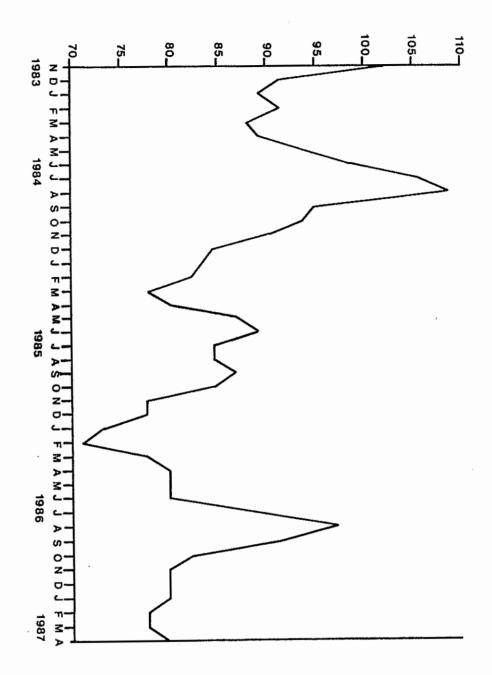
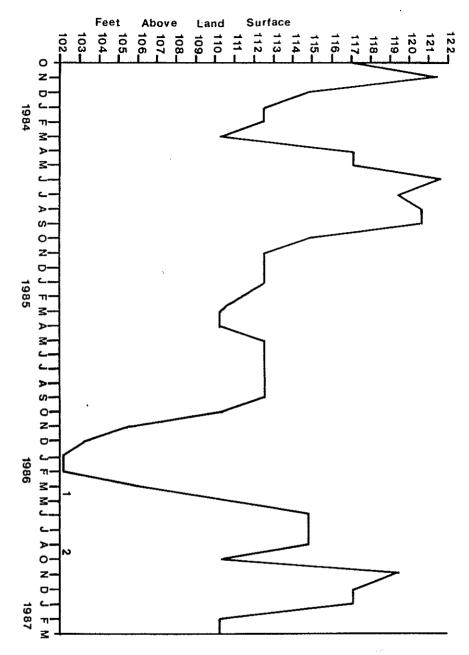
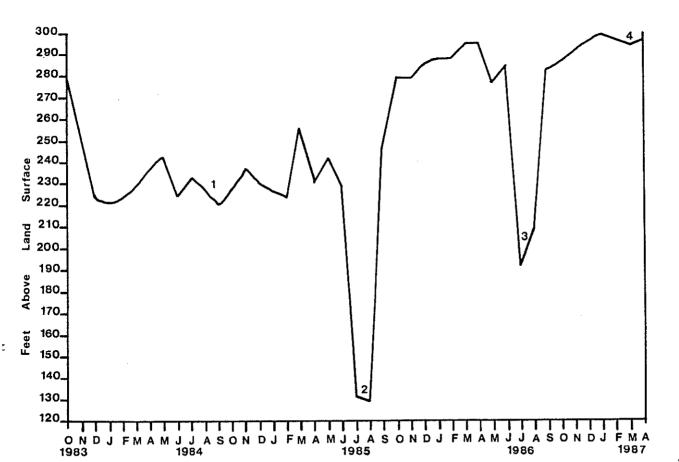


Figure 6
Well #3 Hydrograph
T9S, R14E, Sec. 14, NENESW



- 1.
- No reading for April, 1986 No reading for September, 1986 2.

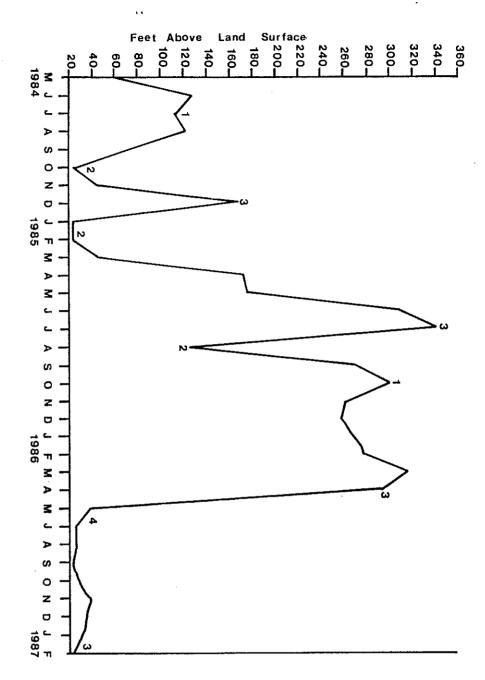
Figure 7 Well #4 Hydrograph T9S, R15E, Sec. 12, SWSWNE



Irrigation
 Irrigation
 Irrigation
 Wells in adjacent area * Figure 8

Well #5 Hydrograph

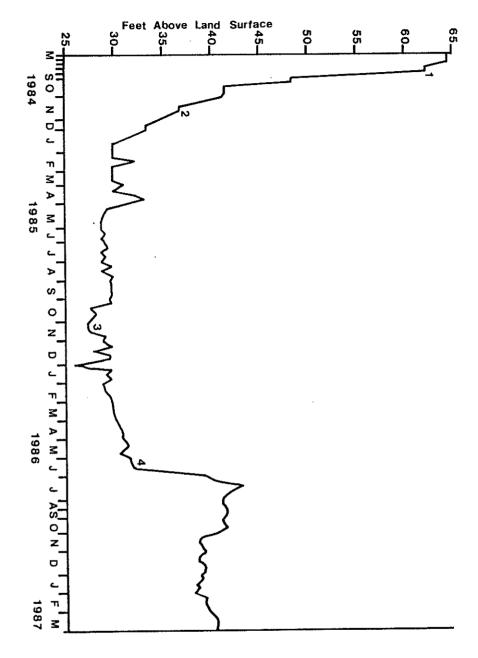
T9S, R14E, Sec. 14, NWSENW



- Small amount flowing 1.
- 2. Completely open
- Shut-in
- Well cleaned out

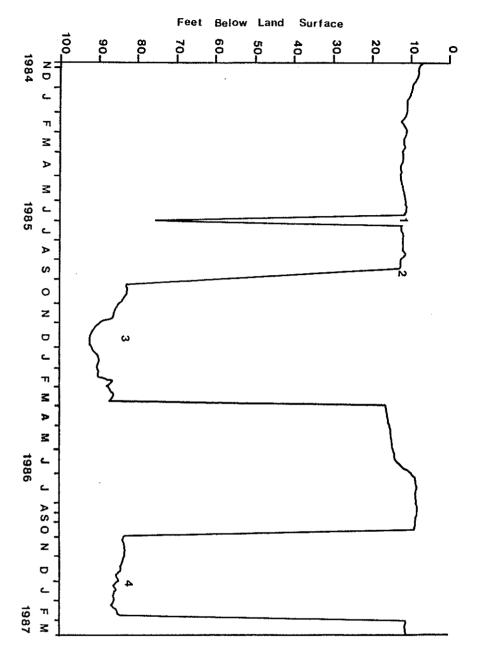
Figure 9 Well #7 Hydrograph TSS, R15E, Sec. 12, SWSWNE

Figure 10
Well #11 Hydrograph
T10S, R17E, Sec. 04, SWNWNE



- Low head power production goes on line from Well #10 Well #13 being developed. 1.
- 2. 3.
- Well #13 goes on line Well #9 repaired

Figure 11 Well #12 Hydrograph T10S, R17E, Sec. 04, SWNENE



- 1.
- Pump test Well being used for heating. Pumping levels Pumping Jevels 2.
- 3.
- 4.

Figure 12 Well #13 Hydrograph T10S, R17E, Sec. 10, NWSENW

Seasonal fluctuations are shown in the hydrographs of wells No.'s 1, 3, and 4 (Figures 4, 6 and 7 respectively). Typically, demand is the greatest during the winter months and recovery occurs in the late summer, early fall. The sharp decline in the months of January and February of 1985 reflects the harsh winter which resulted in increased usage of the geothermal resource.

Well No. 2 (Figure 5) was not used until April, 1984. The record from November, 1983, until that date reflects a shut-in period. The sharp drop in the fall of 1985 represented the time when the owner was testing the system for use in an additional greenhouse. The effect of a new geothermal domestic well drilled in October of 1985 is shown as a depression. From April to the end of June in 1986, the use of the monitored well was expanded to irrigation and domestic consumption because the cold water supply to the area was under repair. This can be seen as a break in the slope of the seasonal recovery.

Well No. 5 (Figure 8) is primarily used for irrigation, hence the deep declines during the summer months. From July to December, 1986, an increase in pressure is shown. This increase was caused by the shut-in of four wells in the adjacent area (T9S, R14E, Sec 14, all 4 wells are shown as No. 6 on Figure 1) that used approximately 4300 gpm from 1983 until late fall, early winter of 1985. Two of the wells were opened and have been utilized since January, 1987 for fish propagation. This subsequent usage is shown by the decline of the water level in the monitored well.

The hydrograph of well No. 7 (Figure 9) is erratic. The well is 1420 feet deep and is cased to 621 feet in poorly consolidated sediments or ash layers. These units continue below the casing where they have the tendency to cave, thus clogging the well. The well was cleaned out during June and July, 1986. Additional casing was not added, therefore the clay continues to clog the well. The well has not been utilized over a continuous

period of time. When the well was flowing freely, the 10 minute shut—in pressures were low, as compared to when the well was completely shut—in over a period of time, the pressures were high. This is evident in the reading of June, 1985. At that time, the pressure had increased to the point where water was leaking from the bolts on the well head. Most of the time the well is either opened completely or is running a small amount; rarely is it completely shut in.

During September, 1984, well No. 10 was opened and used for power production. The well produces at a constant rate of 2470 GPM, and has rarely been shut-in for more than a few hours. production for the power project commenced the response in wells 12 was immediate as indicated in Figures 10 and 11. decline continued until the end of the 1984 heating season. Fluid pressure recovery did not return to pre-production levels. Well 13 was drilled and completed in August, 1984 and was pumped at 280 GPM starting October 1, 1985. Both events show effects on the other two wells. Well 13 was included in the monitoring network in November, 1984 (Figure 12). A pump test was performed 1985, and is shown on the hydrograph. Water levels during the fall and winter months reflect pumping levels. three hydrographs, a dramatic increase in shut-in pressure is evident in July, 1986. At that time, well 9 was rehabilitated with an attendant reduction in leakage.

Well 9 had been flowing for several years at 1500 GPM at the surface and was believed to flow a considerable amount in the subsurface. When drilled in 1970, the flow was estimated to be 2750 GPM and the shut-in pressure was 212 psi. The well was 730 feet deep, completed in rhyolite and was cased to 518 feet in "gray shale and sandstone". There were 300 perforations from 220 to 230 feet and 200 perforations between 460 and 485 feet. The 16 inch steel casing was welded to an 8 inch steel casing at 222 feet. When the well was shut-in during June, 1986, the sound of running water was heard, the well remained warm to the touch, and

response was recorded in wells 11 and 12. All of the above were indications that the well was leaking in the subsurface. was thought that over the years the pressurized water had eroded formations around the bottom of the casing and at the perfoaddition, perforations cut where the 8" rated intervals. the 16 inch casing may have resulted in casing was welded to partial failure of that joint. Cement was pumped into the well to seal the eroded intervals. The well was then redrilled to 743 Casing was set to 640 feet and was pressure-grouted. soon as the well was controlled at depth, the thermal wells in the vicinity immediately responded. The responses can be seen in the hydrographs of wells 11, 12 and 13 (Figures 10, 11 and 12).

Another well, No. 15, was drilled and completed in December, 1986 and the effects of development are shown in the hydrographs of wells 11, 12 and 13 (Figures 10, 11 and 12). Present usage is less than 10 GPM from this new well. Thus this well has an insignificant effect on the overall system.

Annual Discharge

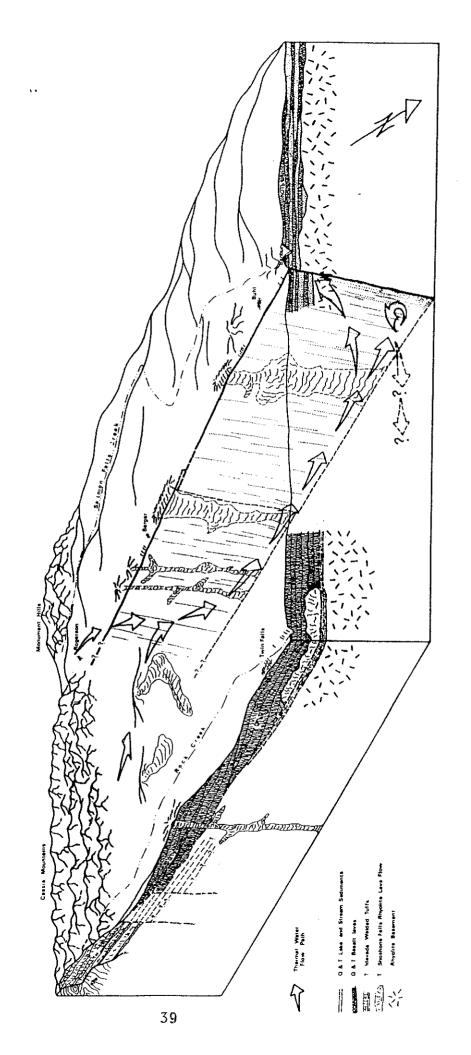
To determine the annual discharge from the thermal aquifer, water rights files and field exam reports were reviewed to determine the maximum permitted usage. Based on discussions with well owners pertaining to seasonal and daily consumption, adjustments were made to the maximum permitted amount. Current annual withdrawal of the aquifer within the Twin Falls - Banbury Study Area was estimated to be 23,600 acre feet per year, (4,364 acre feet for the immediate Twin Falls vicinity and 19,326 for the Banbury area).

PRELIMINARY CONCEPTUAL MODEL OF THE THERMAL SYSTEM

In general, it appears that there are two main directions of flow movement converging in the Twin Falls area: from south to north and from east to west (Young, 1987, personal communication). From the Rock Creek area water follows a generally west-northwest flow path and from the Rogerson area flow is generally to the north. Both flow components encounter the Buhl - Berger Structure Zone in a poorly defined zone southeast of Buhl.

Permeability of the Idavada, both along the flow path and within reservoir rocks, results from fractures related to tectonic movement, sheeted joints and cooling fractures developed during emplacement, intergranular porosity of the non-welded ash flows and air fall tuffs, and voids left between successive flows. Fractures in the BBSZ facilitate the deep circulation of and heating of the water by the regionally high temperature gradient. The heated water is then pushed to the surface in the Banbury area to the northwest by the continued inflow of colder more dense water from the southeast.

It is the authors' opinion based on well testing, similarity of monitoring results, responses to changes in discharge and water chemistry that there appear to be no barriers to thermal water movement within or between the Twin Falls and Banbury portions of the system. While the Twin Falls and Banbury portions of the system are hydrologically connected the source of the heat component at Twin Falls is not clear. The study area is within a of regionally high heat flow which extends from northern Nevada to Yellowstone Mational Park. Heat flow over the zone is generally about 2.5 Heat Flow Units (HFU) (Brott, 1976; Smith, 1980). Lewis and Young (1987, in press) determined heat flow values within the study area to be between 2.2 and 2.5 HFU based on data from thermal gradients derived from bottom hole tempera-The thermal anomaly over the region is beture evaluations. lieved to be related to the Cordilleran Thermal Tectonic Anomaly of Eaton and others (1976) and local thinning of the crust related to Basin and Range extension (Mabey, 1983, p. 13). Heat could be conducted through the aquifer from the Banbury area at a greater rate than the inflow of colder water from the southeast. There could also be inflow of thermal water from other systems from the east and northeast, Figure 13.



Conceptual Model of the Twin Falls - Banbury Thermal System Figure 13.

SUMMARY AND RECOMMENDATIONS FOR FURTHER STUDIES

The geothermal resource in the Twin Falls - Banbury area is characterized by temperatures between 30°C and 70°C (86°F to 158°F) and is confined to the Tertiary Idavada rhyolitic rocks. The probable recharge area is the Cassia Mountains where water could circulate to depths via the regional dip of volcanic strata and structure zones and be heated by the regional thermal gradient.

Increased utilization for heating, irrigation, low-head hydro-power production and fish propagation has led to substantial declines in water levels. Monitoring of the aquifer has shown that temperatures remained constant while water levels are still declining and have not reached equilibrium. The seasonal fluctuations indicate response to the decrease in discharge, not necessarily to recharge. The monitoring also demonstrated that the aquifer responds rapidly to the development and usage of new wells or to the repair and shut-in of existing wells, thus indicating good hydraulic interconnection within the Banbury and Twin Falls portions of the system.

This study was initiated to provide baseline data and to develop a preliminary geologic model of the thermal system. During the course of this investigation, several topics for further study were identified:

- Continued monitoring of the water levels and temperatures of the Twin Falls - Banbury system to expand understanding of the nature and magnitude of declines;
- 2. Additional chemical analyses of thermal waters to determine if changes have occurred over time as a result of withdrawals.
- Comparison of regional thermal water chemistries to determine if there are relationships between the Twin Falls - Banbury and adjacent systems.

- 4. Regional and detailed geologic mapping to: further define the Buhl Berger Structure Zone, determine if other structures exist that may be related to thermal systems and determine if there are other geologic relationships or boundaries between the Twin Falls Banbury and neighboring systems.
- 5. Selected geophysical studies to determine structures at depth and extent of thermal system.

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APPENDIX A

DRILLER'S LOGS

		Depth
	Thickness	(feet below
 Material	(feet)	land surface)

T 8S, R 14E, Sec. 30, NESENW (Casing: 8-inch steel 1 to 204 feet)

Elevation: 2906		Water - Y/N
Quaternary Alluvium		
Brown Clay and Gravel	3	0
Heavy Gravel and Brown Clay	62	3 Y
Banbury Basalt		
Grey Basalt	26	65
Grey Clay	4	91
Brown Clay	13	95
Grey Basalt	30	108
Red Clay	9	138
Grey Basalt	32	147 Y
Brown Basalt	12	179
Grey Basalt	35	191
Green Shale w/Grey Basalt	19	226
Green Shale w/Grey Clay	118	245
Grey Basalt	112	363 Y
Idavada Undifferentiated		
Grey Shale (major field)	5	475 Y
Total Depth		480

		Depth
	Thickness	(feet below
<u> </u>	(feet)	land surface)

T 8S, R 14E, Sec. 30, NENWSE (Casing: 8-inch steel 1 to 80 feet; 6-inch steel 1 to 501 feet)

Quaternary Alluvium 12	Elevation: 2902		Water - Y/N
Gravel and Clay 16 16 Y Gravel 16 16 Y Gravel 23 32 Y Banbury Basalt Grey Basalt 9 55 Brown Scoria and Clay 2 64 Black Basalt 19 66 Brown Basalt 9 88 Brown Clay 5 97 Brown Basalt w/Thin Layers of 5 97 Brown Basalt w/Thin Layers of 40 102 Grey Basalt (Very Hard) 17 173 Red Clay w/Thin Layers of Clay 7 190 Grey Clay 8 197 Grey Basalt w/Thin Layers of 22 205 Brown Basalt 20 227 Grey Grey Clay 22 205 Brown Basalt (Very Hard) 10 258 Grey Basalt w/Green Clay Seams 19 268 Y Small Flow, 2 GPM at 268 8 287 Grey Clay and Sand 8 287 Coarse Sand and Layers of Shale 47 295 <t< td=""><td></td><td></td><td></td></t<>			
Gravel and Clay 16 16 Y Gravel 23 32 Y Banbury Basalt 9 55 Brown Scoria and Clay 2 64 Black Basalt 19 66 Brown Basalt 19 66 Brown Basalt 9 88 Brown Clay 5 97 Brown Basalt w/Thin Layers of 5 97 Brown Basalt (Very Hard) 17 173 Red Clay w/Thin Layers of Clay 7 190 Grey Basalt (Very Hard) 17 173 Red Clay w/Thin Layers of 22 205 Grey Clay 22 205 Brown Basalt w/Thin Layers of 20 227 Grey Basalt (Very Hard) 10 25 Grey Basalt (Very Hard) 10 25 Grey Basalt (Very Hard) 10 25 Grey Clay and Sand 8 287 Coarse Sand and Layers of Shale 47 295 Flow Increased to 5 GPM 34 295 Grey Clay 23 342 Green Clay 25 365 Tan Clay 11 409 Grey Clay 22 400 Tan Clay <t< td=""><td>Brown Sandy Clay</td><td>12</td><td>0</td></t<>	Brown Sandy Clay	12	0
## Banbury Basalt ## Search ## Searc	Gravel	4	12 Y
## Banbury Basalt ## Search ## Searc	Gravel and Clay	16	16 Y
Banbury Basalt Grey Basalt	Gravel	23	32 Y
Brown Scoria and Clay			
Brown Scoria and Clay	Grey Basalt	9	55
Black Basalt		2	64
Brown Basalt. 3 85 Grey Basalt. 9 88 Brown Clay. 5 97 Brown Basalt w/Thin Layers of 31 102 Grey Basalt. 31 142 Grey Basalt (Very Hard). 17 173 Red Clay w/Thin Layers of Clay. 7 190 Grey Clay. 8 197 Grey Basalt w/Thin Layers of 22 205 Brown Basalt. 20 227 Grey Basalt. 11 247 Grey Basalt (Very Hard). 10 258 Grey Basalt w/Green Clay Seams, 19 268 Y Small Flow, 2 GPM at 268 8 287 Coarse Sand and Layers of Shale, 47 295 Flow Increased to 5 GPM 342 342 Green Clay. 25 365 Tan Clay. 11 390 Grey Clay. 22 40 Tan Clay. 11 409 Grey Clay. 22 420 Tan Clay. 4 422 Grey Clay. 9		19	
Grey Basalt. 9 88 Brown Clay. 5 97 Brown Basalt w/Thin Layers of 7 102 Grey Basalt. 31 142 Grey Basalt (Very Hard). 17 173 Red Clay w/Thin Layers of Clay. 7 190 Grey Clay. 8 197 Grey Basalt w/Thin Layers of 22 205 Brown Basalt. 20 227 Grey Basalt. 11 247 Grey Basalt (Very Hard) 10 258 Grey Basalt w/Green Clay Seams, 19 268 Y Small Flow, 2 GPM at 268 8 287 Grey Clay and Sand. 8 287 Coarse Sand and Layers of Shale, 47 295 Flow Increased to 5 GPM 25 365 Grey—Brown Clay. 23 342 Green Clay. 25 365 Tan Clay. 11 409 Grey Clay. 22 420 Tan Clay. 4 442 Grey Clay. 9 446 Y Flow pick			
Brown Clay			
Brown Basalt w/Thin Layers of		· ·	
Sticky Clay 40 102 Grey Basaalt 31 142 Grey Basalt (Very Hard) 17 173 Red Clay w/Thin Layers of Clay 7 190 Grey Clay 8 197 Grey Basalt w/Thin Layers of 22 205 Brown Basalt 20 227 Grey Basalt 11 247 Grey Basalt (Very Hard) 10 258 Grey Basalt w/Green Clay Seams 19 268 Y Small Flow, 2 GPM at 268 287 Grey Clay and Sand 8 287 Coarse Sand and Layers of Shale 47 295 Flow Increased to 5 GPM 342 Grey-Brown Clay 23 342 Green Clay 25 365 Tan Clay 11 390 Grey Clay 8 401 Tan Clay 11 409 Grey Clay 22 420 Tan Clay 4 442 Grey Clay 9 446 Y Flow picks up to 30 GPM Then Dwindles to 5 GPM	Brown Basalt w/Thin Lavers of	J	
Grey Basalt (Very Hard)		40	102
Grey Basalt (Very Hard). 17 173 Red Clay w/Thin Layers of Clay. 7 190 Grey Clay. 8 197 Grey Basalt w/Thin Layers of 22 205 Brown Basalt. 20 227 Grey Basalt. 11 247 Grey Basalt (Very Hard). 10 258 Grey Basalt w/Green Clay Seams, 19 268 Y Small Flow, 2 GPM at 268 8 287 Grey Clay and Sand. 8 287 Coarse Sand and Layers of Shale, 47 295 Flow Increased to 5 GPM 23 342 Grey Brown Clay. 23 342 Green Clay. 25 365 Tan Clay. 11 390 Grey Clay. 8 401 Tan Clay. 11 409 Grey Clay. 22 420 Tan Clay. 4 442 Grey Clay. 9 446 Y Flow picks up to 30 GPM Then Dwindles to 5 GPM 4 442 442 Grey Clay. 60 455	Grev Basalt		
Red Clay w/Thin Layers of Clay. 7 190 Grey Clay. 8 197 Grey Basalt w/Thin Layers of Grey Clay. 22 205 Brown Basalt. 20 227 Grey Basalt. 11 247 Grey Basalt (Very Hard). 10 258 Grey Basalt w/Green Clay Seams, 19 268 Y Small Flow, 2 GPM at 268 287 268 Y Grey Clay and Sand. 8 287 Coarse Sand and Layers of Shale, 47 295 Flow Increased to 5 GPM 23 342 Grey-Brown Clay. 23 365 Tan Clay. 11 390 Grey Clay. 8 401 Tan Clay. 8 401 Tan Clay. 8 401 Tan Clay. 8 401 Tan Clay. 9 446 Y Flow picks up to 30 GPM Then Dwindles to 5 GPM 460 455 Grey Clay. 60 455 Grey Clay. 32 515	Grev Bacalt (Very Hard)		
Grey Clay			
Grey Basalt w/Thin Layers of			
Grey Clay. 22 205 Brown Basalt. 20 227 Grey Basalt. 11 247 Grey Basalt (Very Hard). 10 258 Grey Basalt w/Green Clay Seams, 19 268 Y Small Flow, 2 GPM at 268 287 Grey Clay and Sand. 8 287 Coarse Sand and Layers of Shale, 47 295 Flow Increased to 5 GPM 23 342 Grey-Brown Clay. 25 365 Tan Clay. 11 390 Grey Clay. 8 401 Tan Clay. 11 409 Grey Clay. 22 420 Tan Clay. 4 442 Grey Clay. 9 446 Y Flow picks up to 30 GPM Then 9 446 Y Flow picks up to 30 GPM Then 60 455 Grey Clay. 60 455 Grey Clay. 32 515	Gray Bacalt w/Thin Issuers of	O	197
Brown Basalt		2.2	205
Grey Basalt			
Grey Basalt (Very Hard)			
Grey Basalt w/Green Clay Seams, 19 268 Y Small Flow, 2 GPM at 268 8 287 Grey Clay and Sand. 8 287 Coarse Sand and Layers of Shale, 47 295 Flow Increased to 5 GPM 342 Grey-Brown Clay. 23 342 Green Clay. 25 365 Tan Clay. 11 390 Grey Clay. 8 401 Tan Clay. 11 409 Grey Clay. 11 409 Grey Clay. 4 442 Grey Clay. 9 446 Y Flow picks up to 30 GPM Then Dwindles to 5 GPM 60 455 Grey Clay. 60 455 Grey Clay. 32 515			
Small Flow, 2 GPM at 268 Grey Clay and Sand			
Grey Clay and Sand. 8 287 Coarse Sand and Layers of Shale, 47 295 Flow Increased to 5 GPM 342 Grey-Brown Clay. 23 342 Green Clay. 25 365 Tan Clay. 11 390 Grey Clay. 8 401 Tan Clay. 11 409 Grey Clay. 22 420 Tan Clay. 4 442 Grey Clay. 9 446 Y Flow picks up to 30 GPM Then Dwindles to 5 GPM 60 455 Grey Clay. 60 455 Grey Clay. 32 515		19	268 Y
Coarse Sand and Layers of Shale, 47 295 Flow Increased to 5 GPM 23 342 Grey-Brown Clay. 25 365 Tan Clay. 11 390 Grey Clay. 8 401 Tan Clay. 11 409 Grey Clay. 22 420 Tan Clay. 4 442 Grey Clay. 9 446 Y Flow picks up to 30 GPM Then 9 446 Y Light Grey Clay. 60 455 Grey Clay. 32 515		•	2.27
Flow Increased to 5 GPM Grey-Brown Clay	Grey Clay and Sand	-	= - ·
Grey-Brown Clay 23 342 Green Clay 25 365 Tan Clay 11 390 Grey Clay 8 401 Tan Clay 11 409 Grey Clay 22 420 Tan Clay 4 442 Grey Clay 9 446 Y Flow picks up to 30 GPM Then Dwindles to 5 GPM 60 455 Grey Clay 60 455 Grey Clay 32 515	Coarse Sand and Layers of Shale,	47	295
Green Clay. 25 365 Tan Clay. 11 390 Grey Clay. 8 401 Tan Clay. 11 409 Grey Clay. 22 420 Tan Clay. 4 442 Grey Clay. 9 446 Y Flow picks up to 30 GPM Then Dwindles to 5 GPM 60 455 Light Grey Clay. 60 455 Grey Clay. 32 515			
Tan Clay			-
Grey Clay			
Tan Clay			390
Grey Clay		8	- - -
Tan Clay	Tan Clay		409
Grey Clay	Grey Clay	22	420
Flow picks up to 30 GPM Then Dwindles to 5 GPM Light Grey Clay	Tan Clay	4	442
Flow picks up to 30 GPM Then Dwindles to 5 GPM Light Grey Clay	Grey Clay	9	446 Y
Light Grey Clay 60 455 Grey Clay 32 515	Flow picks up to 30 GPM Then		
Light Grey Clay 60 455 Grey Clay 32 515			
Grey Clay 32 515		60	455
- •			
Grey Brown Sand Stone			

Idavada Undifferentiated			
Black Igneous Rock	22	590	Y
Grey-Brown Igneous Rock	21	612	Y
Black Igneous Rock	67	633	
Total Depth		700	

		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 8S, R 14E, Sec. 30, NWNWSE (Casing: 12-inch steel 0 to 35 feet; 8-inch steel 1 to 145 feet)

Elevation: 2900		Water -Y	/N
Quaternary Alluvium			
Top Soil	15	0	
Medium Hard Rock and Gravel	10	15	
Clay and Rock	60	25	
Banbury Basalt			
Medium Hard Black Rock	25	85	Y
Clay and Black Rock	10	110	
Rock and Clay Mix	60	120	
Hard Solid Lava	35	180	
Clay and Water Strips	220	215	Y
Clay and Rock Mix	5	435	Y
Idavada Undifferentiated			
Hard Black Rock	10	440	
Total Depth		450	

**		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 8S, R 14E, Sec. 32, SWSENE (Casing: Not on File)

Elevation: 2950		Water - Y/N
Quaternary Alluvium		
Brown Sandy Clay	3	0
Boulders	27	3
Grey Sand	5	30
Boulders, Cobble Stones & Gravel	11	35
Boulders	6	46 Y
Brown Clay	6	52
Grey Clay & Gravel	6	58
Gravel	12	64
Brown Clay and Gravel	9	76
Banbury Basalt	,	, •
Grey Basalt	18	85
	5	103
Brown Clay	5	108
Grey Basalt	6	
Brown Clay		113
Grey Basalt	38	119
Brown Basalt	4	157
Red Clay	8	161
Black Basalt	26	169
Grey Basalt	15	195
Grey Basalt	4	210
Grey Clay & Basalt	11	214
Grey Clay	16	225
Grey Basalt	43	241
Dark Grey Clay	8	284
Grey Shale w/Layers of Sand	8	292
Grey-Black Basalt	12	300 Y
Dark-Grey Clay	48	312
Black Basalt	15	360
Sticky Clay	15	375
Grey Shale Clay & Rocks	17	390
	19	407
Green Clay		426
Basalt	2	
Grey Clay & Shale	12	428 Y
Idavada Undifferentiated		4.4.0
Black Rhyolite	31	440
Grey Sand Stone	40	471
Grey-Brown Rhyolite	34	511
Brown-Grey Sandstone Shale & Rocks		
Grey-Green Shale, Rock & Clay	15	545 Y
Main Flow		

Grey Rhyolite	9	560 Y
Grey Clay	3	569
Light Grey Rhyolite w/Layers		
of Green Shale	27	572 Y
Light Grey Rhyolite w/Layers		
Shale	24	599
Grey Clay Sticky	6	623
Light Grey Rhyolite	16	629
Total Depth		645

		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 8S, R 14E, Sec. 32, SENENE
(Casing: 12-inch steel 0 to 49 feet;
8-inch steel 1.5 to 447 feet;
6-inch steel 400 to 815 feet)

Elevation: 2906		Water - Y	/N
Quaternary Alluvium	5	0	
Sand	_	0 5	
Sand & Clay	43	5	
Banbury Basalt	7.0	4.0	
Black Basalt	78	48	
Sandstone	24	126	Y
Sandstone	40	150	Y
Sand, Lots of Water	10	190	Y
(Sandstone) Layer of Sand	120	200	Y
Sand	6	320	Y
Sandstone74°	54	326	
Sandstone80°	20	380	Y
Sand74°	11	400	
Rock	11	411	
Sandstone74°	13	422	
Clay74°	12	435	
Banbury Basalt69°	33	447	
Clay & Basalt	10	480	
Sandstone69°	30	490	Y
Sand82°	10	520	-
Sandstone93°	105	530	Y
Sand	2	635	-
Grey Shale	3	637	
Sandstone	30	640	
Idavada Undifferentiated	30	040	
Black or Dark Grey Rhyolite	8	670	
Grey Shale			
-	4	678	
Sandstone	18	682	
Grey Shale	42	700	
Black Rhyolite	4	742	
Grey Shale	14	746	
Sandstone	30	760	
Black Rhyolite	70	790	Y
Green-Blue Shale108°	30	860	
Grey Shale	40	890	
Grey-Black Rhyolite	15	930	
Grey Shale	2	945	
Grey-Black Rhyolite	7	947	
Grey-White Shale	4	954	
Dark Grey Shale w/Some of			
Lighter Grey	62	958	
J	- -		

Grey-Green Shale	10	1,020
Grey Rhyolite w/Layers of Grey Shale.	50	1,030
Grey Shale & Clay	5	1,080
Grey Rhyolite w/Grey Clay	15	1,085
Grey Shale	20	1,100
Grey Rhyolite w/Layers of Clay	30	1,120
Grey Shale	30	1,150
Grey Shale w/White Soft Clay	2	1,180
Grey Shale	38	1,182
Grey/Green Shale	60	1,220
Grey Shale w/Clay Layers	20	1,280
Grey Shale w/Thin Clay Layers	80	1,300
Total Depth		1,380

		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 8S, R 14E, Sec. 30, NENESW
(Casing: 16-inch steel 1 to 23 feet;
12-inch steel 1 to 47 feet;
6-inch steel 1 to 444 feet)

Elevation: 2940		Water-Y/N
Quaternary Alluvium		
Tan Clay	15	0
Dark Grey Clay	9	15
Dark Grey Clay w/ ?	30	24 Y
Tertiary Basalt		
Black Basalt (Hard)	20	54
Grey Shale	6	74
Red-Brown Clay	6	80
Black Basalt	18	86
Red-Brown Silty Clay	7	104 Y
Red & Black Cinders	11	111 Y
Grey Clay	84	122
Red Clay w/Rock Layers	16	206
Grey Basalt w/Clay	18	222
Grey Basalt (Hard)	18	240
Grey-Brown Clay	8	258
Grey-Brown Basalt	6	266
Grey Basalt	24	272
Grey Shale	11	296
Grey Basalt	2	307
Grey Shale w/Stick Clay	17	307
Light Tan Shale	11	
Grey Clay & Shale	27	326 337
	3	
Dark Grey Basalt		364 Y
Grey Green Shale	18	367
Grey Shale	7	385
Green Shale	7	392
Grey Clay	33	399
Tan Clay	9	432
Grey-Green Clay w/Rock Layers	9	441
Grey-Green Shale w/Rock Layers	12	450
Idavada Undifferentiated	_	
Grey Rhyolite	7	462 Y
Grey Shale w/Rock Layer	15	469 Y
Black Rhyolite	40	484 Y
Grey Clay	4	524
Grey Shale w/Rock Layers	37	528 Y
Black Rhyolite	22	565 Y
Grey Shale w/Clay Layers	13	587 Y
Decomposed Brown Rhyolite	10	600 Y
(Major Flow)		
Total Depth		610

•		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 8S, R 14E, Sec. 32, NENWSE (Casing: 10-inch steel 1 to 299 feet)

Elevation: 2931		Water	- Y/N
Quaternary Alluvium			
Brown Sandy Clay	6	0	
Boulders & Gravel	61	6	
Boulders	16	67	
Brown Clay	2	83	
Tertiary Basalt			
Grey Basalt (Hard)	48	85	
Grey Basalt	5	133	
Red Clay	6	138	
Grey Basalt w/Clay Layers	16	144	
Grey Basalt (Hard)	32	160	
Brown Clay	11	192	
Grey-Green Clay or Shale	14	203	
Black Basalt	26	217	Y
Grey Clay, Shale & Basalt	8	243	•
Grey Basalt	16	251	Y
Grey Brown Shale	11	267	Y
Black Basalt w/Shale Layers	5	278	1
Gravel	10	283	
Idavada Undifferentiated	10	203	
	22	203	
Green Shale & Clay		293	
Grey Rhyolite	21	315	
Grey Rhyolite & Shale Layered	27	336	
Grey-Brown Rhyolite	46	363	
Grey Rhyolite	4	409	
Grey Rhyolite (at 435' Small Flow)	105	413	Y
Light Shale w/Rhyolite Layered	5	518	
Grey Rhyolite	59	523	Y
Grey Rhyolite w/Clay Layers	30	582	Y
Grey Rhyolite	26	612	
Grey Rhyolite w/Clay Layers	13	638	
Green Shale w/Layers of Rhyolite	12	651	
Brown Clay (Sticky)	6	663	
Black-Brown Rhyolite	16	669	
Grey Rhyolite	11	685	
Grey Rhyolite w/Green Shale	23	696	
Black Rhyolite	35	719	
Green Cinders	4	754	Y
Black Rhyolite	15	758	-
Grey Rhyolite	5	773	Y
Grey Rhyolite	4	778	-
Brown Rhyolite	24	782	Y
Grey Rhyolite	26	806	Y
Grey-Brown Physlite	16		_
Grey-Brown Rhyolite	Τ0	832	Y

Grey Rhyolite	178	848	
Black Rhyolite w/clay seams	13	1,026	
Grey Clay	5	1,039	
Black Rhyolite	8	1,044	
Brown Clay (Sticky)	14	1,052	Y
Grey Rhyolite w/Layers Sticky Clay	9	1,066	
Grey Rhyolite	5	1,075	
Total Depth		1,080	

•		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 8S, R 14E, Sec. 32, NWNWSE (Casing: 12-inch steel 1 to 39 feet; 8-inch steel 2 to 591 feet)

Elevation: 2975		Water - Y/N
Quaternary Alluvium		
Top Soil	22	0
Gravel	4	22
Clay	8	26
Banbury Basalt		
Red Lava, (16" to 38')	67	34
Grey Lava	124	101
Lava & Clay	6	225 Y
Grey Lava	127	231
Grey Lava, Strips Blue Clay	167	358
Blue Clay80°	12	525 Y
Grey Lava	35	537
Blue Clay86°	17	572 Y
Grey Lava (12" to 591')	35	589
Idavada Undifferentiated		
Broken Grey Rhyolite110°	2	624 Y
Hard Red-Brown Rhyolite	64	626
Hard Grey Rhyolite	183	690
Broken Grey Rhyolite120°	4	873 Y
Hard Grey Rhyolite	25	877
Soft Grey Rhyolite	93	902
Broken Grey Rhyolite140°	15	995 Y
Total Depth		1,010

		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 8S, R 14E, Sec. 33, NWNENW (Casing: 6-inch steel 1 to 361 feet)

Elevation: 2900		Water - Y/N
Quaternary Alluvium		
Clay	31	0
Gravel	2	31
Clay	92	33
Sand	15	125
Banbury Basalt		
Black Basalt	33	140 Y
Clay	47	173
Sand	25	220 Y
Clay	55	225
Black Basalt	35	280
Conglomerate	4	315
Black Basalt (Medium to Hard)	36	319
Black Basalt (Very Hard)	100	355
Blue Clay	117	455
Blue Shale	86	572
Broken Blue Shale	12	658
Blue Clay	145	670 Y
Blue Shale (Hard)	120	815
Blue Clay	214	935
Blue Shale	53	1,149
Blue Clay	40	1,202
Idavada Undifferentiated		·
Reddish Brown Rhyolite	33	1,242 Y
Total Depth		1,275

,		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 8S, R 14E, Sec. 33, NWNWSW

(Casing: 10-inch steel 1 to 199 feet;
8-inch steel 148.5 to 497.5 feet;
6-inch steel liner 494 to 694 feet;
Liner is perforated:
40 perforations - 497 to 505 feet
32 perforations - 674 to 682 feet)

Elevation: 2918			Vater -	- Y/N
Quaternary Alluvium	4.5		•	
Sandy Clay	13		0	
Clay w/Light Grey Sand	11		13	
Sand w/Clay	19		24	Y
Brown Sand w/Small Gravel	5		43	Y
Boulders w/Sand	4 67		48 52	Y
Grey Clay	07		34	
Banbury Basalt	3		119	
Black Basalt (Hard)			122	
Basalt, Broken w/Clay	17 5		139	
Basalt, Broken w/Clay	20		144	
Basalt, Broken w/Clay	20		164	
	2		166	
Grey Clay	10		168	
Brown Clay	2		178	
Black Basalt, Fractured w/Clay Tan Clay	9		180	
Blue-Grey Clay	3		189	
Blue Shale (First Artesian Flow)	7		192	Y
bide Shale (First Arcesian Flow)	•	Flows Grouted		I
Blue Shale	3	riows Grouced	199	
Black Basalt (Hard)	1		202	
Blue Shale	10		202	
Black Basalt (Hard)	8		213	
Brown Shale	2		221	
Basalt, Broken w/Clay	12		223	
Tan Clay	5		235	
Green Clay	11		240	
Green Shale	13		251	
Grey Clay	17		264	
Idavada Undifferentiated	- '		201	
Black & Green Rhyolite	4		281	
Light Brown Shale	4		285	
Dark Green Shale	11		289	
Light Grey Clay	30		300	
Green Clay	4		330	
Shale & Green Clay	37		334	
Grey Shale	13		371	
Green Shale	18		384	

Grey Shale	24	402	Y
Brown Clay	14	426	
Brown Basalt	4	440	
Shale & Green Clay	8	444	
Blue Shale	24	452	
Grey Shale	12	476	
Dark Blue Shale	14	488	
Black Basalt	8	502	
Red-Brown Rhyolite	9	510	
Black Rhyolite	3	519	
Brown Rhyolite	5	522	
Black Rhyolite	9	527	
Sandstone	4	536	
Black Rhyolite	7	540	
Decomposed Rhyolite, Alternating			
- · · · · · · · · · · · · · · · · · · ·	111	547	
Black Rhyolite (Hard)	3	658	
Grey Clay	2	661	
Rhyolite (Very Hard)	31	663	
Total Depth		694	

		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 8S, R 14E, Sec. 33, SENWSW (Casing: 12-inch steel 1 to 119 feet; 8-inch steel 1.5 to 278 feet; 6-inch steel 185 to 470 feet)

Elevation: 2907		Water - Y/N
Quaternary Alluvium		
Brown Clay	8	0
Brown Sand	27	8 Y
Tertiary Basalt		
Grey Basalt	13	35
Grey-Brown Clay	19	48
Grey Basalt	5	67
Brown Clay	20	72
Grey Basalt	45	92
Brown Basalt	5	137
Grey Basalt	26	142
Dark Brown Clay	7	168
Grey Basalt	34	175
Brown Clay & Brown Basalt	4	209
Brown Basalt	6	213
Brown Basalt	34	219
Green Clay	4	253
Tan Clay w/Thin Layers of Basalt	13	257
Tan Clay w/Shale Layers	36	270
Hard Grey Shale	21	306 Y
Grey-Brown Clay w/Shale in Layers	25	327 Y
Grey-Brown Shale w/Clay		
in 1' to 2' Layers	26	352
Grey Clay	31	378
Light Tan Clay	7	409
Dark Green Shale	4	416
Light Green Clay	21	420
Grey-Brown Shale	9	441
Grey Shale	20	450
Dark Grey Shale	17	470
Light Grey Shale	20	487
Idavada Undifferentiated		
Black Rhyolite	3	507 Y
Brown Rhyolite	46	510 Y
Black Rhyolite	22	556 Y
Grey Shale	6	578 Y
Black Rhyolite	48	584 Y
Black Rhyolite w/Thin Layers of		
Sticky Clay	50	632 Y
Brown Rhyolite	61	682 Y
Grey-Brown Rhyolite	57	743 Y
Total Depth		800
_		

-		Depth
	Thickness	(feet below
<u>Material</u>	(feet)	land surface)

T 8S, R 14E, Sec. 33, NWNWSW (Casing: 8-inch steel 1 to 259 feet) 6-inch steel 171 to 482 feet)

Elevation: 2912		Water - Y/N
Quaternary Alluvium		_
Brown Sandy Clay	33	0
Brown Clay	3	33
Brown Sandy Clay	20	36
Gravel & Sandy Clay	17	56 Y
Grey Sand	39	73 Y
Gravel & Boulders	30	112 Y
Grey Sandy Clay	3	142 Y
Banbury Basalt		
Grey Basalt	3	145
Gravel	8	148
Brown Clay	5	156
Gravel & Clay	25	161
Grey Basalt	42	186
Brown Basalt & Clay	11	228
Grey Basalt w/Clay	3	239
Grey Basalt	29	242
Grey Shale	14	271
Brown Clay	16	285
Grey-Green Shale & Clay	57	301 Y
Green Shale w/Sand & Gravel	21	358
Gravel	5	379
Green Shale	9	384
Green Clay	24	393
Grey Clay	31	417
Dark Grey Shale	6	448
Grey Clay	16	454
Grey Shale	48	470
Shale & Sand in Thin Layers	11	518 Y
Idavada Undifferentiated		-
Black Rhyolite	53	529
Grey Shale	5	582
Brown Rhyolite	53	587
Grey Shale	14	640
Black Rhyolite	14	654 Y
Black Rhyolite	19	668 Y
Brown Rhyolite	13	687 Y
Total Depth	T)	700
Total Depth		700

		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 9S, R 14E, Sec. 4, NWNWNW (Casing: 8-inch steel 0 to 95 feet; 6-inch steel 1 to 215 feet)

Elevation: 2959	<u></u>	Water - Y/N
Tertiary Basalt		
Sand & Dirt	35	0
Decomposed Lava (Medium-Hard)	30	35
Clay & Gravel (Water ?)	75	65 Y
Idavada Undifferentiated		
Black Rock (Hard)	110	140
Black Rock (Medium-Hard)	100	250
Grey Rhyolite (Hard)	75	350 Y
Broken Spot		
Grey Rhyolite (Hard)	155	425
Broken Rhyolite	5	580 Y
Black Rock (Hard) w/Clay Seams	55	585
Black Rock (Hard)	260	640
Total Depth		900

,		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 9S, R 14E, Sec. 4, NESESW (Casing: 8-inch steel 1 to 41 feet; 6-inch steel 1 to 380 feet)

Elevation: 2996		Water - Y/N
Tertiary Basalts		
Dirt & Rock	6	0
Brown Basalt	14	6
Grey Basalt	70	20
Black Basalt & Clay	5	90
Black Basalt	25	95
Brown Clay	11	120
Grey Silt	42	131
Undifferentiated Idavada Volcanics		
Grey Shale (Hard)	2	173
Grey Sandy Clay	50	175
Grey Clay	23	225 Y
Grey Shale (Hard)	3	248
Grey Clay	19	251
Grey Shale	110	270
Brown Shale	27	380
Grey Shale	91	407
Black Rhyolite	104	498 Y
Grey Rhyolite	248	602
Total Depth		850 Y

		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 9S, R 15E, Sec. 12, SWNESW (Casing: 8-inch steel 2 to 110 feet; 6-inch steel 30 to 621 feet)

Elevation: 3065		Water - Y/N
Tertiary Basalts		
Top Soil	3	0
Boulders & Gravel	7	3
Grey Lava	123	10
Brown Clay	8	133
Blue Clay	19	141 Y
Brown Sandstone	44	160 Y
Idavada Pyroclastics		
Red Lava	4	204 Y
Black Lava	60	208 Y
Brown Lava & Red Ash	11	268 Y
Black Lava	49	279 Y
Grey Sand & Blue Clay	39	328 Y
Black Lava	39	367 Y
Reddish Brown Lava Ash	14	406 Y
Black Lava	50	420 Y
Brown Lava	130	470 Y
Reddish Brown Lava Ash	50	600 Y
Sandstone & Clay	50	650 Y
Grey Lava (Hard)	50	700 Y
Blue Clay	10	750 Y
Grey Lava	75	760 Y
Soft Blue Clay	203	835 Y
Hard Rock	37	1,038 Y
Broken Rock	45	1,075 Y
Clay	155	1,120 Y
Clay	9	1,275 Y
Hard Black Rock	16	1,284 Y
Black Rock, Layers, Broken	120	1,300 Y
Total Depth		1,420

•		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 9S, R 17E, Sec. 29, NWNWSE (Casing: 6-inch steel 0 to 640 feet) Well has been reconstructed.

Elevation: 3154		Water - Y/N
Tertiary Basalts		
Brown Sand	7	0
Tan Sandy Clay	5	7
Boulders in Gravel	26	13
Loose Basalt	12	39 Y
Basalt Boulders	7	51
Cinders & Clay	2	53 Y
Basalt Boulders	10	60
Cinders & Clay	11	70 Y
Basalt Boulders	6	81
Basalt (Soft)	37	87
Grey Basalt	24	124
Brown Basalt	16	148
Grey Basalt	29	164
Brown Clay & Rock	9	193
Brown Basalt	17	202
Grey Basalt	34	219 Y
Brown Clay w/Rock	20	253
Grey Basalt	39	273
Brown Clay	9	312
Grey Basalt	40	321
Grey Clay w/Rock	21	361
Grey Basalt	6	382
Brown Basalt	8	388
Grey Basalt	38	396
	10	434
Grey Clay	15	444
Grey Basalt Lake Sediments	15	यु यु यु
	8	459
Tan Clay		
Grey Sandy Clay w/Layered Shale	11	467
Grey Sand	2	478 Y
Idavada Pyroclastics	_	400
Tan Shale	6	480
Grey Shale & Sandstone	32	486
Consolidated Black Shale or Rock(?)	5	518
Grey Shale	24	523 Y
Brown Rhyolite w/Layers of ?	58	547 Y
Green Rhyolite (Soft)	25	605 Y
Green Rhyolite (Hard)	20	630 Y
Brown Rhyolite (Very Hard)	35	650 Y
Pink Rhyolite (Very Hard)	15	685 Y

Black & Pink Rhyolite (More Water)	2	700	Y
Black & Pink Rhyolite (More Water	3	702	Y
125 to 175 GPM)	5	705	Y
Pink & Black Rhyolite	15	710	Y
Brown & Pink Rhyolite	18	725	
Total Depth		743	Y

		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 9S, R 17E, Sec. 32, NESESE (Casing: 12-inch steel 1 to 71 feet; 8-inch steel 2 to 493 feet)

Elevation: 3637		Water -	Y/N
Quaternary/Tertiary Basalts		-	
Top Soil	5	0	
Lava Boulders	5	5	
Grey Lava	23	10	
Cements (?). Gravel	11	33	
Grey Lava	9	46	
Brown Sandy Clay	4	55	
Grey Lava	12	59	
Lava Cinders & Gravel	22	71	Y
Brown Lava Cinders	7	93	
Brown Lava Cinders & Clay	40	100	
Grey Lava (Hard)	40	140	
Brown Clay	6	180	
Grey Lava	20	186	
Reddish Brown Lava	49	206	
Brown Sandy Clay w/Gravel	24	255	
Tan Sticky Clay	39	279	
Grey Lava (?)	20	318	
Brown Sandy Clay	17	338	
Grey Lava	37	355	
Brown Sandy Clay	63	392	
Blue Sticky Clay	39	455	Y
Grey Lava (Hard)	75	494	
Reddish Brown Clay	8	569	
Brown Clay	48	577	
Grey Sandy Clay	29	625	
Brown Sandy Clay	29	654	
Idavada Pyroclastics			
Grey Rhyolite (Hard)	63	683	
Reddish Brown (?)	14	746	
Grey Rhyolite (Very Hard)	30	760	
Black Rhyolite (Hard)	15	790	
Rhyolite (Slightly Softer)	15	805	
Brown Lava & Clay	20	820	
Black Rhyolite (Hard)	20	840	
Decomposed Lava	20	860	
Broken Black Rhyolite (Hard)	80	880	
Sandstone	12	960	
Sandstone (Very Soft)	2	972	
Sandstone, Brown Shale	53	974	
Black Lava	33	1,027	
Black Rhyolite (Very Hard)	10	1,060	
Rhyolite (Slighter Softer)	30	1,070	

Softer Form Again	7	1,100	
Black Rhyolite (Very Hard)	8	1,107	
Red & Black Rhyolite	35	1,115	
Red & Black Rhyolite (Broken)	15	1,150	
Medium-Hard Strips of Soft Broken			
Rhyolite	40	1,165	
Soft White Shale	20	1,205	Y
Rhyolite, Broken (Medium-Hard)	10	1,225	Y
Red Rhyolite, Broken (Soft)	5	1,235	
Red Rhyolite, Broken (Hard)	40	1,240	
Total Depth		1,280	

•		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 9S, R 17E, Sec. 33, NWNWNW

(Casing: 24-inch steel 1 to 26 feet; 16-inch steel 1 to 128 feet; 12-inch steel 1 to 261 feet; 8-inch steel 1 to 592 feet)

Elevation: 3174		Water - Y/N
Bonneville Flood Deposits		
Sand & Dirt	8	0
Grey Lava Boulders & Sand	18	8
Grey Lava Boulders & Sand	30	26
Grey Rhyolite Boulders & Cinder Mix	23	56
Tertiary Basalts		
Grey Basalt (Hard)	7	79
Grey Rhyolite & Cinders (Soft)	30	86 Y
Grey Lava & Some Clay	7	116 Y
Blue Clay & Lava (Some Caving)	17	123
Grey Lava (Hard)	17	140
Grey & Brown Clay/Grey Rhyolite	10	157
Grey Rhyolite w/Layers of Clay	36	167
Grey Rhyolite & Clay	23	203
Grey Basalt (Hard)	34	226
Grey Basalt (Very Hard)	14	260
Brown Basalt (Hard)	6	274
Black Basalt (Hard)	32	280
Broken Basalt (Crevice)	2	312
Black Basalt - Some Brown	1	314
Black Basalt - Big or Broken		
Boulders	14.5	315
Grey Basalt (Very Hard) Andesite?	47.5	329.5
Black Basalt (Andesite?), Broken	12	377
Black Basalt? Andesite?	5	389
Black Basalt, Broken	19	394
Grey Basalt (Very Hard)	25	413
Broken Basalt (Softer) - Trace of		
Cold Flowing Water	14	438
Solid Basalt	8	452
Lake Sediments		
Layers of Sandstone & Clays - Some		
Warmer Water	92	460 Y
Layers of Clay & Rock - Water		
Increasing - About 150 GPM Temp.		
Increase from 80° to 92° F.	40	552 Y

Idavada Pyroclastics			
Broken Rhyolite w/Shale Layers - Water			
Increased about 600 GPM			
Temp. 99°	78	592	Y
Broken Rhyolite w/Shale Layers -Big			
Increases in Water Temp.,			
Increased 102°-103°	80	670	Y
Total Depth		750	

		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 10s, R 13E, Sec. 20, NESENE (Casing: 20-inch steel 0 to 25 feet; 16-inch steel 0 to 100 feet; 12-inch steel 1 to 255 feet)

Elevation: 3428		Water - Y/N
Idavada Pyroclastics		
Boulders	25	0 Y
Brown Rhyolite	37	25
Red Clay	26	62
Loose Rock	12	88 Y
Brown Rhyolite	48	100
Loose Rock & Clay	107	148
Brown Rhyolite	145	255
Grey Rhyolite (Hard)	62	400
Brown Rhyolite	18	462 Y
Grey Rhyolite	178	480
Blue Clay (Sticky)	37	658
Brown Rhyolite	35	695 Y
Brown & Grey Loose Rhyolite	120	730 Y
Grey Rhyolite	70	850
Red Rhyolite (Loose)	55	920 Y
Grey Rhyolite (Hard)	65	975
Grey Brown (Loose)	45	1,040 Y
Red Brown (Loose)	20	1,085 Y
Grey Rhyolite (Hard)	135	1,105
Loose	27	1,240 Y
Hard Rhyolite	88	1,267
Rhyolite (Broken)	35	1,355 Y
Rhyolite (Hard)	70	1,390
Total Depth		1,460

		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 10S, R 17E, Sec. 4, NESESW (Casing: 16-inch steel 1 to 367 feet; 12-inch steel 2 to 1611 feet;

Elevation: 3668		Water - Y/N
Quaternary Basalts		
Brown Clay	9	0
Black Basalt	11	9
Brown Silt	4	20
Black Basalt	11	24 Y
Brown Silt	8	35
Broken Basalt	4	43 Y
Black Basalt	32	47
Brown Clay	3	79
Black Basalt	21	82
Grey Basalt	17	103
Brown Basalt	33	120
Grey Basalt	9	153
Brown Basalt	25	162
Grey Basalt (Very Hard)	40	187
Brown Clay	5	227
Grey Basalt	3	232
Brown Clay	9	235
Gravel & Brown Clay	39	244
Grey Clay	22	283
Brown Clay w/Some Gravel	30	305
Grey Basalt	3	335
Grey Clay	5	338
Grey Basalt	74	343
Grey Basalt (Creviced)	8	417
Grey Clay	8	425
	12	433
Grey Basalt (Creviced & Hard)	50	445
Grey Basalt (Hard)	68	495
Grey Basalt w/Clay Layers		
Grey Basalt (Very Hard)	7	563
Grey Basalt w/Clay Seams	13	570
Grey Basalt (Hard)	2	583
Grey Clay	10	585
Grey Basalt	15	595
Grey Basalt (Hard)	15	610
Brown Clay	8	625
Shosĥone Falls Rhyolite		
Grey Andesite (Very Hard)	30	633
Grey Andesite (Very Hard)	15	663
Brown Clay	9	678
Grey Basalt w/Clay Layers	11	687
Grey Clay	7	698
Brown Clay & Rocks	28	705

Grey Basalt	17 35 75 10 31 2 4 49 7 15 3 11 19 1	733 750 785 860 870 901 903 907 956 963 978 981 992 1,011 1,012	
Brown Clay	10 15 5 20 15 20 2 3 13	1,020 1,030 1,045 1,050 1,070 1,085 1,105 1,107 1,110	Y
Brown Rhyolite	30 15 45	1,125 1,155 1,170	Y Y
Red-Brown Rhyolite	335 105 85 60	1,215 1,550 1,655 1,740	
Grey Rhyolite	20 17 45 25 21 34 91 43 9 16 66 33	1,800 1,820 1,837 1,882 1,907 1,928 1,962 2,053 2,096 2,105 2,121 2,187 2,220	Y

		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 10s, R 17E, Sec. 4, SENWSW

(Casing: 16-inch steel 1 to 512 feet;
12-inch steel 1 to 1248 feet;
10-inch steel 1184 to 1345 feet;
8-inch steel 1293 to 1453)

(Perforations: 16 rows per foot 1293 to 1333; 12 rows per foot 1333 to 1353; 8 rows per foot 1353 to 1453; Size of perforations: 1/4" x 3")

Elevation: 3662		Water - Y/N
Quaternary and Tertiary Basalt		
Brown Clay	15	0
Clay & Broken Rock	7	15 Y
Grey Basalt	23	22 Y
Brown Clay	4	45
Grey Basalt	14	49
Clay	13	63
Grey Basalt	9	76
Clay	13	85
Grey-Brown Basalt	32	98 Y
Grey Basalt	34	130
Brown Basalt	17	164
Grey Basalt	37	181
Talc	2	218
Grey Basalt	21	220
Brown Sandy Clay w/Gravel	137	241
Soft Brown Conglomerate	8	378
Grey Basalt	30	386
Brown Clay	32	416
Grey Clay	14	448
Brown Clay	37	462
Grey Clay	9	499
Grey Basalt (Hard)	27	508
Brown Clay	8	535
Brown Sandy Clay	97	543
Grey Sandy Clay	28	640
Grey Sandy Clay	46	668
Grey Sandy Clay	8	714
Grey Basalt	26	722 ·
Reddish Brown Clay	8	748
Grey Basalt	22	756
Brown Basalt w/Layers of Sticky Clay.	5	778
Brown Clay	5	783

Grey Basalt (Hard)	65	788
Brown Basalt	5	853
Grey Basalt	38	858
Grey Clay (Sticky)	16	896
Tertiary Basalt		
Grey Basalt (Very Hard)	76	912
Grey Clay	2	988
Grey Basalt	5	990
Grey Clay (Sticky)	28	995
Grey Basalt	5	1,023
Clay & Shale w/Fine Gravel	34	1,028
Black Basalt	8	1,062
Brown Clay (Sticky)	12	1,070
Grey Clay & Shale (Sticky)	8	1,082
Grey Brown Shale	21	1,090
Grey Basalt,	2	1,111
Grey Clay	1	1,113
Grey Basalt	8	1,114
Grey Clay & Rock Layered	22	1,122
Black Basalt	13	1,144
Grey Clay & Shale	19	1,157
Grey Basalt	5	1,176
Grey Shale	5	1,181
Grey Basalt w/Thin Clay Layers	12	1,186
Grey Basalt	4	1,198
Grey Basalt w/Thin Clay Layers	7	1,202
Brown Clay	6	1,209
Grey-Brown Basalt	15	1,215
Grey Basalt (Hard)	3	1,230
Grey-Brown Basalt	15	1,233
Lake Sediments		
Grey Andesite (Very Hard)	19	1,248
Brown Shale (Caved In)	48	1,267
Brown Shale	6	1,315
Layers of Sand, Shale & Clay	14	1,321
Light Brown Clay	15	1,335
Brown Rhyolite Chips & Clay	7	1,350
Idavada Pyroclastics		
Red-Brown Rhyolite	123	1,357
(Broken 1409-1455)		
1455 on Broken, Creviced and Caving		
Formation		
Total Depth		1,480
Main Flow - 1422' to 1455'		

		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 10S, R 17E, Sec. 5, NENESE (Casing: 12-inch steel 1 to 18 feet; 8-inch steel 2 to 1091 feet; 6-inch steel 1080 to 1200 feet)

Elevation: 3647		Water -	Y/N
Quaternary and Tertiary Basalts			
Top Soil	3	0	
Hard Pan	13	3	
Grey Lava (16" to 19')	12	16	
Brown Ash	5	28	
Brown Lava	18	33	
Brown Cinder	5	51	
Brown Lava (54°)	84	56	Y
Red Lava	14	140	
Brown Lava	32	154	
Grey Lava (54°)	68	186	
Brown Sandy Clay & Gravel	124	254	
Grey Lava (54°)	26	378	
Brown Clay (55°)	23	404	
Blue Clay	105	427	
Grey Lava & Rock	13	532	
White Clay	3	545	
Brown Sandstone	6	548	
Grey Lava	104	554	
Brown Sandstone	12	658	Y
Grey Lava (Very Hard)	55	670	
Lake Sediments			
Brown Clay	10	725	
Blue Clay Shale	14	735	
Idavada Pyroclastics			
Black Rhyolite (63°)	166	749	
Black Lava (Softer)	23	915	
Black Lava (Hard) (63°)	63	938	
Decomposed Lava or Rhyolite	74	1,001	
Brown & Red Sandstone, Broken		·	
Rhyolite (85°)	10	1,075	Y
Red Rhyolite (12" to 1091')	35	1,085	
Broken Red Rhyolite (90°)	3	1,120	Y
Red Rhyolite	27	1,123	
Red Rhyolite Sand (90°)	15	1,150	Y
Broken Red Rhyolite	15	1,165	-
Red Rhyolite (8" to 1200')	69	1,180	Y
Broken Rhyolite (91°)	3	1,249	Ÿ

Solid Red Rhyolite	48	1,252	
Broken Red Rhyolite	5	1,300	
Solid Red Rhyolite	70	1,305	
Broken Red Rhyolite	7	1,375	
Solid Red Rhyolite	68	1,382	Y
Total Depth		1,450	
Started Flowing 20 GPM at 1,075'			
Main Flow From 1180' to 1.252'			

		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 10S, R 17E, Sec. 10, NWSENW (Casing: 16-inch steel 1 to 23 feet; 12-inch steel 2 to 900 feet)

Elevation: 3717		Water - Y/N
Quaternary Basalt		
Top soil	12	0
Grey Lava	36	12
Lava Ash	5	48
Grey Lava	77	53
Red Lava	7	130
Brown Lava	8	137
Red Lava	5	165
Grey Lava	68	170
Brown Lava	20	238
Brown Lava	5	258
Shoshone Falls Rhyolilte		
Brown Rhyolite	11	263
Grey Rhyolite	41	274
Grey Rhyolite	123	315
Softer Grey Rhyolite (More Water)	20	438
Red & Grey Broken Rhyolite - Water	13	458
Red & Grey Broken Rhyolite - Lots		
of Water	47	471
Brown Rhyolite	10	518
Broken Grey Rhyolite - Water	₂ 3	528
Hard Grey Rhyolite	26	531
Broken Softer Red Rhyolite	26	557
Solid Grey Rhyolite	4	583
Grey Rhyolite (Very Hard)	19	587
Broken Brown Rhyolite - Water	6	606
Solid Grey Rhyolite	13	612
Broken Grey Rhyolite - More Water	13	625
Grey Rhyolite (Hard)	60	638
Grey Shale (Some Black Rock) -		
More Water	16	698
Black Rhyolite, Broken - Water	83	714
Lake Sediments	0.5	,
Broken Brown Rock	6	797
Brown Clay	6	803
Green Clay & Shale	33	809
Idavada Pyroclastics	J J	009
Broken Black Rock	28	842
Soft Red Rhyolite	9	870
Medium Black Rock	31	879
MEGIUM BIAUK RUCK	ЭT	013

Hard Black Rock	16	910
Soft Brown Rock	6	926
Hard Black Rock	59	932
Soft Brown Sandstone	64	991
Grey Decomposed Rhyolite - Hit		
81° Water	30	1,055
Grey Decomposed Rhyolite	36	1,085
Red Rhyolite	125	1,121
Broken Red Rhyolite - 89°	38	1,246
Solid Brown Rhyolite	124	1,284
Red Rhyolite	8	1,408
Broken Green & Pink Rhyolite -		
Some Water	4	1,416
Solid Brown Rhyolite	85	1,420
Grey Rhyolite	195	1,505
Total Depth		1,700

Depth Thickness (feet below Material (feet) land surface)

T 10S, R 17E, Sec. 14, SESWSW (Casing: 6-inch steel 1 to 580 feet)

Elevation: 3786		Water - Y/N
Quaternary Basalt		_
Soil	10	0
Hard Pan	2	10
Lava	37	12
Lava Red	2	49
Lava	10	51
Clay & Shale, Medium Flow	12	61
Lava	4	73
Lava (Hard)	14	77
Lava Red, Medium Flow	3	91
Lava (Hard)	11	94
Lava	16	105
Loose Formation - Small Flow	3	121
Lava	12	124
Lava Red	5	136
Lava	50	141
Lava Red - Small Flow	14	191
Lava	60	205
Piller Falls Mud Flow		
Boulders - Small Flow	28	265
Shoshone Falls Rhyolite	20	200
Lava	17	293
Reaming well from 6 1/4 to 8 1/4	1	253
Lava (Hard)	92	310
Crevice & Talc	1	402
	9	403
Lava (Hard) - More Water	4	412
Clay	_	
Broken Formation of Hard Rock	2	416
Lava	2	418
Clay	2	420
Lava (Hard)	36	422
Clay	2	458
Lava (Hard)	15	460
Lava (Soft)	3	475
Lava (Hard)	22	478
Lava (Soft)	10	500
Lava (Hard)	68	510
Clay	4	578
Blue Clay	6	582
Lava	12	588
Red Lava - More Water	5	600
Black Lava	22	605
Clay	2	627
Lava	9	629

Clay	3	638
Lava	7	. 641
Lava (Soft) - More Water	3	648
Lava (Hard)	21	651
Blue Mud & Clay - More Water	6	672
Lava	9	678
Blue Mud & Boulders	23	687
Rock	28	710
Blue Clay - More Water	1	738
Lava (Hard)	21	739
Lava (Hard) - Reaming Out Wall	70	760
Lava (Hard)	100	830
Lava (Hard) - More Water	5	930
Lava (Hard)	45	935
Lake Bed		
Limestone	45	980
Quartz Crystals	6	1,025
Shale	24	1,031
Lake Bed	30	1,055
Idavada Pyroclastics		
Red Granite	46	1,085
Sandstone - Water Strong		1,131
Sandstone - Water Strong	19	1,135
Total Depth		1,154

•		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 10S, R 18E, Sec. 6, NWNWNW

(Casing: 10-inch steel 1 1/2 to 6 1/2 feet; 8-inch steel 1 to 149 feet; 6-inch steel to 1173 feet)

Elevation: 3585		Water - Y/N
Quaternary Basalts		
Brown Clay	1	0
Quaternary Basalt (Hansen Butte)	29	1
Light Clay	6	30
Brown Basalt	14	36
Grey Basalt	14	50
Brown Clay	2	64
Red-Brown Basalt	6	66
Grey Basalt	38	72
Grey Silty Clay and Gravel	15	110
Brown Basalt	5	125
Basalt (Slightly Green)	5	130
Gravel, Silt & Sand	15	135
Shoshone Falls Rhyolite		
Decomposed Rhyolite	17	150
Brown Rhyolite	202	167
Grey-Brown Rhyolite	16	369
Grey Rhyolite	60	385
Grey-Brown Rhyolite	20	445
Brown Rhyolite	27	465
Grey Rhyolite	11	492
Dark Sand	4	503
Black Rock (Soft & Shiny)	11	507
Lake Sediments		
Rock and Clay in Layers	19	518
Brown Clay	9	537
Idavada Pyroclastics		
Brown Rhyolite	4	546
Brown Clay	3	550
Grey Rhyolite	45	553
Red-Brown Clay	4	598
Black Rhyolite	24	602
Red-Brown Clay	5	626
Brown Rhyolite	4	631
Tan and Green Shale	21	635
Andesite (Very Hard)	89	656
Black Rhyolite & Grey Clay in Layers.	34	745
Brown Clay (Sticky)	4	779
Andesite & Grey Clay Layered	-	
(Very Hard)	97	783
Grey Clay & Shale	13	880

Y

		Depth
	Thickness	(feet below
Material	(feet)	land surface)

T 10S, R 19E, Sec. 1, SESESE

(Casing: 8-inch steel 0 to 47 feet;
6-inch steel 0 to 1460 feet;
5-inch steel 13 to 1925 feet;
3-inch steel 1900 to 2160 feet

Perforations: 50 - 1400 to 1925 feet;
150 - 2100 to 2160 feet)

Elevation: 3952		Water - Y/N
Quaternary Basalt		
Brown Clay	42	0
Grey Basalt	10	42
Grey Scoria	31	52
Grey Basalt	32	83
Grey Scoria	21	115
Brown Cinders	34	136
Caving Rock	10	170
Grey Basalt	20	180
Black Basalt	30	200
Brown Scoria	20	230
Grey Scoria	18	250
Grey Basalt (Hard)	6	268
Grey Scoria	23	274 Y
Grey Basalt	11	297 Y
Grey Scoria	66	308 Y
Grey Basalt	6	374 Y
Grey Scoria	23	380 Y
Grey Basalt & Clay	12	403 Y
Brown Basalt	50	415 Y
Grey Basalt	17	465 Y
Brown Clay	4	482
Brown Basalt	36	486
Brown Clay	3	522
Brown Scoria	3	525
Brown Basalt w/Clay Layers	13	528
Brown Clay	8	541
Brown Basalt	18	549
Brown Clay w/Layers of Basalt	7	567
Brown Basalt	31	57 <i>4</i>
Grey-Brown Basalt	36	605
Red Clay & Brown Basalt	23	641 Y
Grey-Brown Basalt	16	664
Grey Basalt	63	680
Grey Sand	4	743
Grey Basalt	17	747
Brown Basalt	13	764
DIOWH Dasait	т Э	1 O 4

Idavada Pyroclastics		
Grey-Brown Rhyolite	27	777
Brown Rhyolite	92	804
Grey Rhyolite	215	896
Black Rhyolite	11	1,111
Brown Shale	11	1,122
Brown Rhyolite	9	1,133
Brown Clay	3	1,142
Brown Rhyolite	7	1,145
Grey Rhyolite	13	1,152
Andesite (Very Hard)	9	1,165
Red Clay	2	1,103
Grey Rhyolite	7	1,176
Brown Clay	19	1,183
Grey Rhyolite	12	1,202
Brown Clay	3	1,214
Grey Clay & Rhyolite	14	· · · · · · · · · · · · · · · · · · ·
Cross Physical	2	1,217
Grey Rhyolite		1,231
Grey Clay	42	1,233
Grey Rhyolite	11	1,275
Grey Clay & Shale	16	1,286
Grey Rhyolite	59	1,302
Grey Clay (Sticky)	7	1,361
Grey Sand	5	1,368
Dark Grey Clay (Sticky)	40	1,373
Grey Sandy Clay	27	1,413
Tan Clay (Sticky)	6	1,440
Grey Clay (Sticky)	14	1,446
Andesite (Very Hard)	69	1,460
Brown Clay (Sticky)	4	1,529
Sandy Clay	5	1,533 Y?
Red-Brown Sandy Clay	17	1,538
Grey Clay (Sticky)	20	1,555
Grey Shale	4	1,575
Grey Shale (Sticky)	11	1,579
Grey Clay (Sticky)	5	1,590
Grey Sandy Clay	5	1,595 Y?
Grey Shale	18	1,600
Grey Clay & Shale	10	1,618
Black Rhyolite	42	1,628
Red Clay	15	1,670
Grey Shale & Clay Layered	15	1,685
Grey Rhyolilte	5	1,700
Grey Shale	10	1,705
Grey Rhyolite	5	1,715
Grey Clay (Sticky)	115	1,720
Light Grey Pumice Clay	8	1,835
Light Green Pumice Clay	10	1,843
Light Grey Clay (Sticky)	22	1,853
Grey Shale (Caving)	75	1,835
Dark Grey Clay w/Thin Shale Layers	75 39	· · · · · · · · · · · · · · · · · · ·
Grey Shale	2	1,950
Dark Grey Clay w/Thin Layers of Shale	21	1,989
	33	1,991
Light Blue-Grey Clay	33	2,012

Dark Grey Clay	12	2,045	
Dark Grey Shale	6	2,057	
Light Grey Clay	49	2,063	
Light Grey Sand	7	2,112	
Grey Shale w/Alternating Layers of		0.110	
Green Sand	37	2,119	¥
Grey-Green Rock (Rock is fine grained	4	2 1 5 6	
and Very Hard)	4	2,156	
Total Depth		2,160	

·		Depth		
	Thickness	(feet below		
Material	(feet)	land surface)		

T 14S, R 15E, Sec. 16, SESWSE (Casing: 8-inch steel 2 to 143 feet)

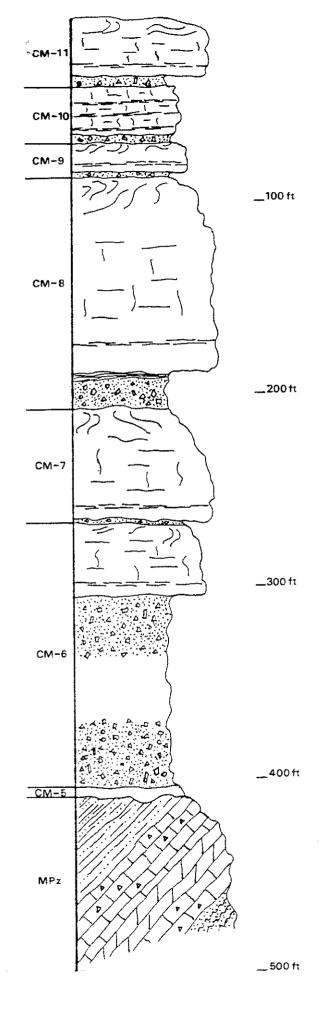
Elevation: 4931		Water - Y/N
Tertiary Basalts	_	
Top Soil	5	0
Grey Lava	22	5
Red-Brown Sandy Clay	32	27
Grey Lava	56	59
Red-Brown Lava Ash & Cinders	48	115
Brown Clay & Gravel	72	163
Brown Sandy Clay	45	235
Black Lava Cinders	32	280
Black Lava (Hard)	11	312
Idavada Pyroclastics		
Brown Rhyolite	147	323
Broken Brown Rhyolite	65	470
Red Cinders & Talc	33	535
Grey Rhyolite	7	568
Broken Red Rhyolite	43	575
Hard Red Rhyolite	17	618
Broken Red Rhyolite	57	635
Brown Clay	12	692
Broken Brown Rhyolite & Grey Clay	112	704
Solid Red Rhyolite	109	816
Solid Grey Rhyolite	51	925
Tan Clay	24	976
Broken Red Rhyolite & Clay	60	1,000
Hard Red Rhyolite	22	1,060
Red Clay	56	1,082
Hard Red Rhyolite	95	1,138
Broken Red Rhyolite	19	1,233
Black Lava	8	1,252
Red Lava Ash	15	1,260
Brown Clay & Lava	32	1,275
Black Sandy Clay	6	1,307
Brown Rhyolite	152	
Grey Rhyolite	45	1,313 Y 1,465
Red Rhyolite	30	•
Void	11	1,510
Crevise and Broken Rock	10	1,540
Hard Rock	10 49	1,551
	4 7	1,561
Total Depth		1,610

APPENDIX B

GENERALIZED STRATIGRAPHIC SECTION

FOR THE

CASSIA MOUNTAINS



Stratigraphic Section for the Cassia Mountains APPENDIX B Generalized

samples were taken from densely each includes a densely welded rhyolitic ash-flow All geochemistry and the underlying non-welded pyroclastics. Units identified as CM-5 through CM-11 welded units. tuff

MPz - Pre-Cenozoic marine sedimentary rocks.

APPENDIX C

CHEMICAL ANALYSIS OF IDAVADA PYROCLASTICS

AND THE SHOSHONE FALLS RHYOLITE

TECHNIQUES AND RESULTS

CHEMICAL ANALYSES TABLES

Samples CM10-1, CM10-2, CM9-1, CM9-2, CM11-3, CM11-2, CM11-1, CM10-4 and CM10-3 were analyzed at the University of Utah Research Institute Earth Science Laboratory. The samples were crushed and split in a tungsten carbide shatterbox, pulverized to -200 mesh. The samples were then fused with lithium borate followed by the appropriate dilutions. Major elements, except silica were then analyzed by Inductively Coupled Spectrometry. Silica and the trace elements were analyzed by inductively coupled plasma spectrometry.

Samples CM8-1, CM8-2, CM8-3, CM8-4, CM8-5, CM7, CM5, SF1, SF2 and WC1 were analyzed at Rice University Department of Geology and Geophysics. The samples were crushed and prepared for analysis in the form of lithium metaborate fusion followed by appropriate dilutions. Six of the samples were prepared in duplicate. Major element concentrations were analyzed using Inductively Coupled Spectrometry.

	CM 11-3	CM 11-2	CM 11-1	CM 10-4	CM 10-3
% oxide					
0 011244					
sio ₂	70.70	73.05	66.80	65.87	73.41
Al ₂ O ₃	13.00	12.24	14.87	11.83	11.71
TiO ₂	0.53	0.49	0.63	0.61	0.47
Fe as Fe ₂ O ₃	4.24	3.96	4.86	4.23	3.21
MnO	0.08	0.07	0.10	0.07	0.05
CaO .	1.61	1.19	1.87	4.74	1.06
MgO	0.32	0.29	0.97	0.50	0.24
к ₂ 0	5.27	2.91	3.88	4.10	4.00
Na ₂ O	3.18	3.32	2.11	2.86	2.61
P205	0.08	0.08	0.61	0.11	0.06
Total	99.01	97.60	96.70	94.92	96.82
LOI	1.48	. 44	2.87	3.00	1.62
Ba (% oxide)	0.14	0 14	0.11	0.12	0.14
Sr Sr	103.	95.	149.	143.	68.
	34.			8.	13.
Co Cu	6.	12. 10.	7.	10.	6.
Pb	11.	14.	24.	11.	
Zn	105.	93.	20.	105.	18.
Sb			103. 33.		75. 31.
Li	20.	9.	34.	12.	16.
Be	3.	3.	4.	2.	3.
Zr	545.	462.	470.	312.	480.
La	98.	91.	92.	87.	93.
Ce	173.	160.	163.	143.	167.
F	670.	90.	540.	340.	310.

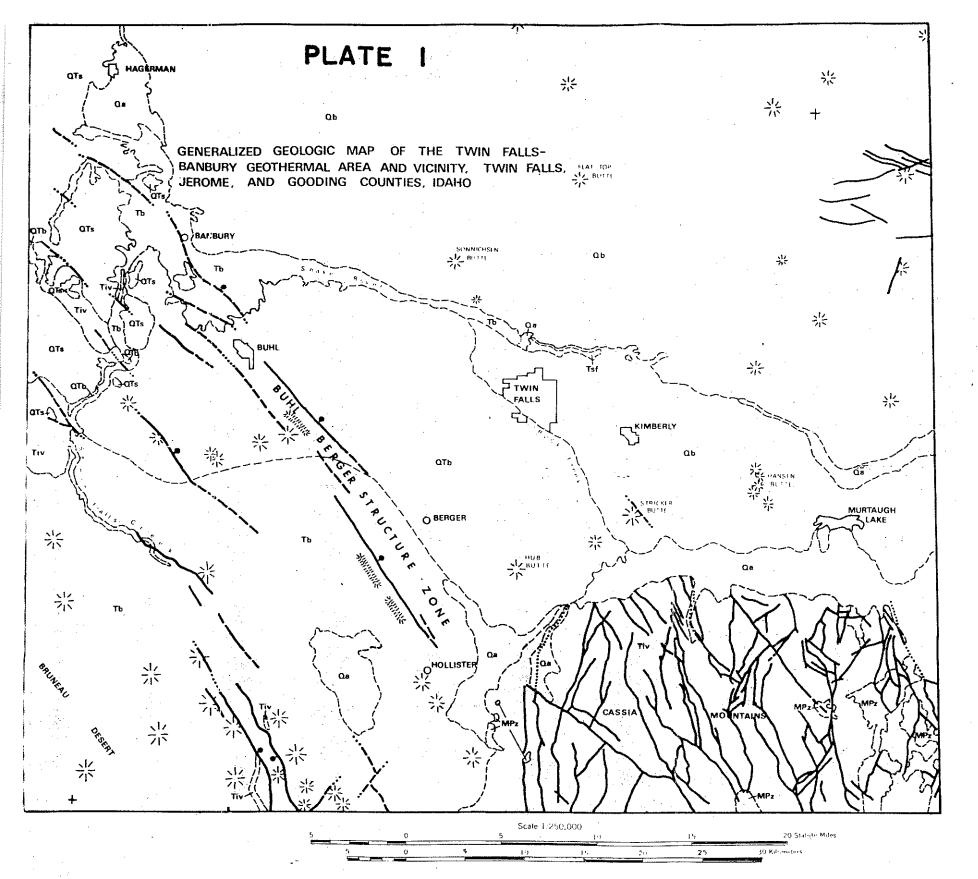
Minor elements in parts per million.

	CM 8-5	CM 8-4	CM 8-3	CM 8-2	CM 8-1
% oxide					
sio ₂	73.70	74.62	73.76	71.81	73.21
Al ₂ 0 ₃	11.84	11.64	11.90	12.71	12.21
TiO ₂	0.31	0.30	0.28	0.64	0.49
Fe as Fe ₂ O ₃	2.42	2.46	2.23	4.56	2.93
MnO .	0.04	0.02	0.02	0.07	0.03
CaO	0.59	0.46	0.27	1.70	0.68
MgO	0.13	0.07	0.06	0.49	0.17
к ₂ 0	5.66	5.28	5.34	4.38	5.91
Na ₂ O	2.72	3.23	3.07	3.28	2.10
P2 ^O 5	0.07	0.07	0.05	0.15	0.10
Total	97.48	98.15	96.98	99.79	97.83
LOI	2.55	0.41	0.37	0.56	3.01

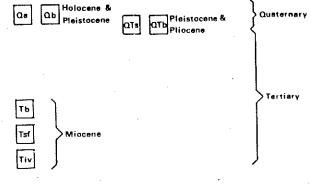
	CM 10-2	CM 10-1	CM 9-2	CM 9-1
% oxide				
sio ₂	72.30	72.41	71.40	68.80
Al ₂ O ₃	11.39	11.67	12.20	12.63
TiO ₂	0.51	0.48	0.47	0.56
Fe as Fe ₂ O ₃	3.20	3.25	4.03	4.73
MnO	0.05	0.05	0.07	0.08
CaO .	1.28	1.10	1.38	1.87
MgO	0.42	0.24	0.27	0.39
к ₂ 0	4.79	5.32	3.35	4.95
Na ₂ O	2.33	2.69	3.02	2.96
P ₂ O ₅	0.03	0.06	0.08	0.11
Total	96.30	97.27	96.27	97.08
LOI	2.62	1.50	1.70	1.28
Ba (% oxide)	0.12	0.14	0.13	0.14
Sr	88.	68.	82.	112.
Со	33.	41.	11.	32.
Cu	6.	6.	6.	7.
Pb	15.	18.	18.	L10.
Zn	65.	93.	93.	103.
Sb	L30.	L30.	33.	L30.
Li	17.	16.	18.	9.
Ве	3.	3.	3.	3.
Zr	436.	481.	567.	554.
La	81.	92.	95.	94.
Ce	141.	164.	167.	163.
F	210.	310.	540.	120.

Minor elements in parts per million.

	CM 7	CM 5	SF 1	SF 2	WC 1
% oxide					
SiO ₂	65.21	71.93	71.54	68.65	74.72
A1203	13.04	12.02	13.72	13.48	12.35
TiO ₂	0.77	0.43	0.69	0.69	0.36
Fe as Fe ₂ O ₃	5.34	3.57	4.66	4.69	2.95
MnO .	0.08	0.04	0.07	0.07	0.05
CaO	2.45	0.87	1.82	1.98	0.61
MgO	0.74	0.20	0.72	0.67	0.17
к20	4.77	5.54	4.71	4.93	4.92
Na ₂ O	2.88	2.44	3.53	3.17	3.39
P205	0.22	0.08	0.18	0.16	0.06
Total	95.50	97.12	101.64	98.49	99.58
LOI	1.87	2.95	.36	1.75	.26



Correlation of Map Units



Description of Map Units

Mesozoic &

Paleozoic

- Qs Quaternary Sedimentary Rocks Holocene and Pliestocene gravels, sands, silts and clays. Includes Meion Gravel along the Snake River in the western part of the area and Snake River deposits in the east. Also includes large areas of stream and lake deposits along the northern margin of the Cassie Mountains.
- Quaternary Olivine Basalt Lave Flows Includes flows from Ransen Butte and flows of Snake River Group Basalt from ounerous vent areas north of the Snake River Canyon. Many flows show young constructional features such as as and pahoehoe lavs surfaces and pressure ridges.
- QTs Lover Plioceme to Upper Pleistocene Continental Sedimentary Units Includes stream and lake rediments primerily of the Clenns Petry Formation. Composed of gravel, sand, silt, clsy and interbedded volcanic ash beds.
- QTb Lower Pliocene to Upper Pliestocene Olivine Bassit Lava Flows Includes flows from Hub Butte and other vents in the southwestern and western part of the area. Some units within the Glenns Ferry Formation are included. Constructional features generally removed by erosion or obscured by Lorsa.
- The lower Miocene to Upper Plincene Banbury Basalt Consists of lavs flows of olivine basalt locally interbedded with stress and lake sediments.

 Basalts are commonly eltered to greenish brown saprolite with residual spheroids of undecomposed rock. Sediments composed of lenticular channel deposits of sand and pebble gravel and light colored silt, clay and distomite in massive lake deposits.
- Taf Rhyolite Lava Flow of Shosnone Falls Porphyritic light gray devitrified single lavs flow. Contains zones of sheeted or platy fractures. Columnar jointing isn't procounced but strong vertical fractures are abundant. Phenocrysta are predominantly plaginclose with minor amounts of pyroxene and opaque oxides.
- Tiv Welded Ash-Flow Tuff Sheety of the Idavada Volcanic Group Predominantly densely welded units separated by sirfall, water-lain and
 non-welded ash-flow tuff. Densely welded units are typically zoned from
 base to top: bedded base surge clastics, basal vitrophere, thick
 massive central devitrified lithoidal zone, and a thin upper lithoidal
 vapor phase zone showing prominent flow structures and scattered
 vitrophere.
- MPz Hesozoic and Paleozoic Marioe Sedimentary Rocks of the Cassia Mountains - Includes limestone, dolonitic limestone, eiltstone, obert, and quartrite. Formations include lower Trissic Diawoody, lower Permian Phosphoria and Grandeur Toogue of the Park City and several other locally delimested Permian and possible Ordovician units.

Fault - dashed where inferred or approximately located; dotted where concealed; bar and ball on downthrown side

Geologic contact (all contacts are generalized)

쑮

Basaltic vent

郑阳比

Fissure cruption vent